

# An Introduction to Granular Matter

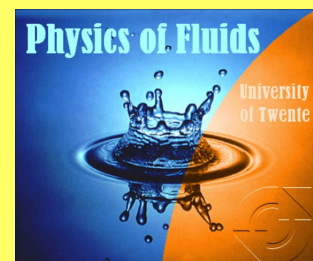
D. Van der Meer & D. Lohse



J.M. Burgerscentrum



**University of Twente**  
*Enschede - The Netherlands*



# GRANULAR MATTER is everywhere:

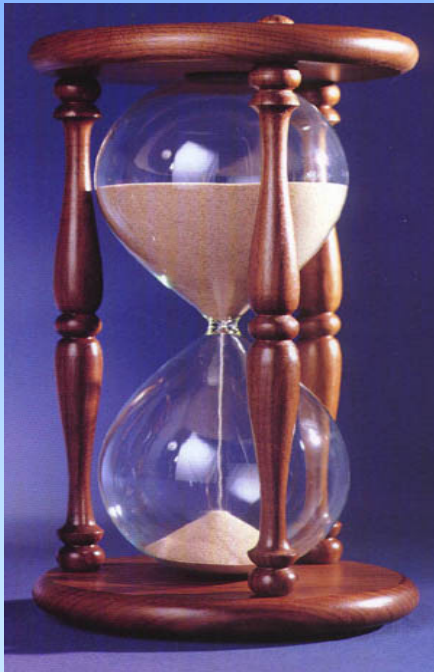


- in nature: beach, soil, snow, desert, mountains, sea floor, Saturn's rings, asteroids ...
- in industry: mining, pharmaceutical, food, construction, chemical ...

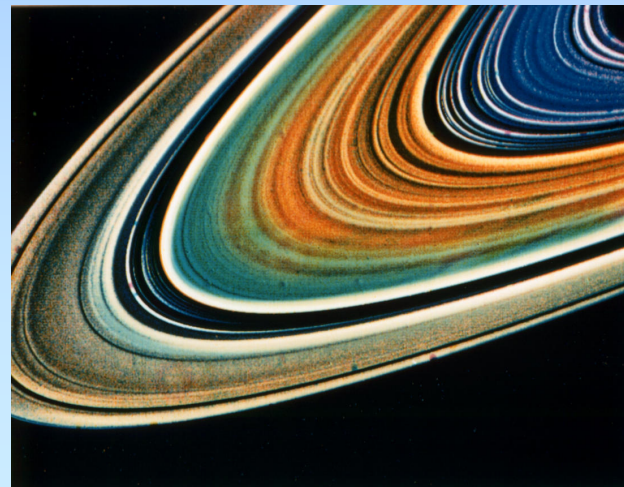


# Granular Matter can behave like...

... a solid



... a liquid



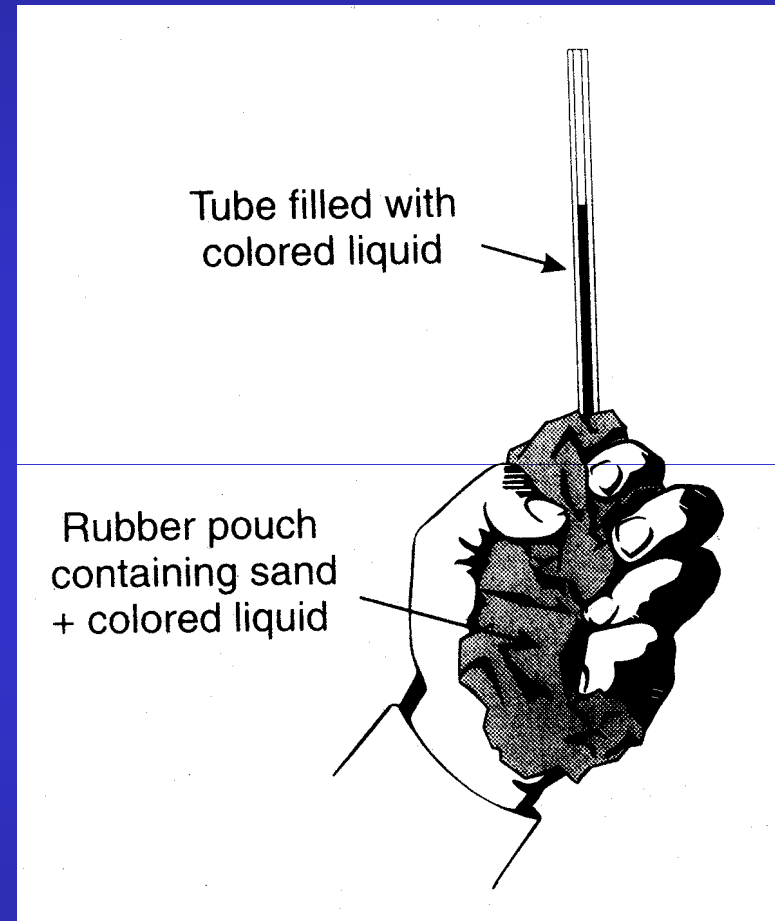
... or a gas



When solid,  
Granular Matter is a special solid



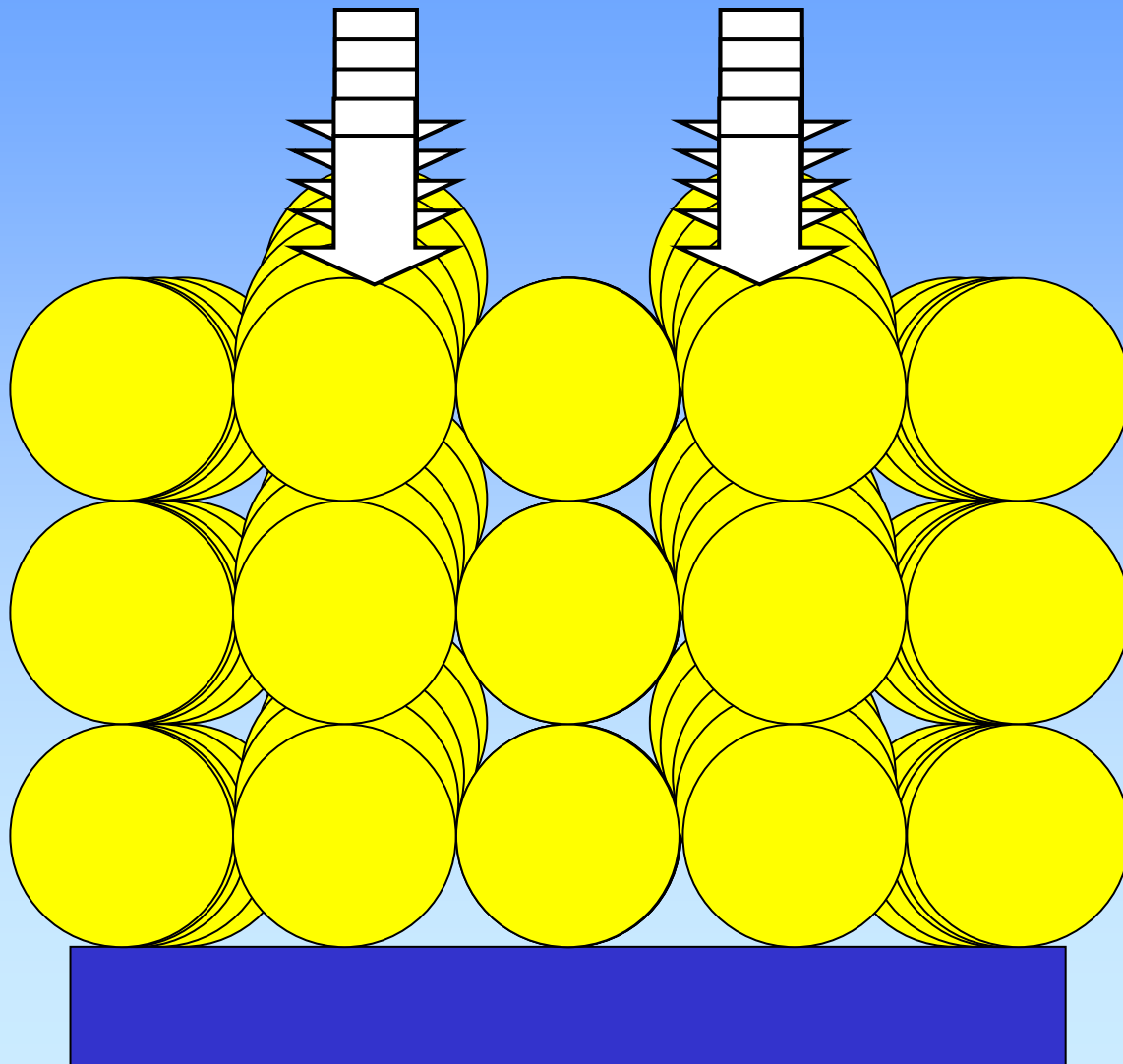
# Reynolds dilatancy



Osborne Reynolds (1885):

“A strongly compacted granular medium **dilates** under pressure”.

# What causes the dilatancy ?



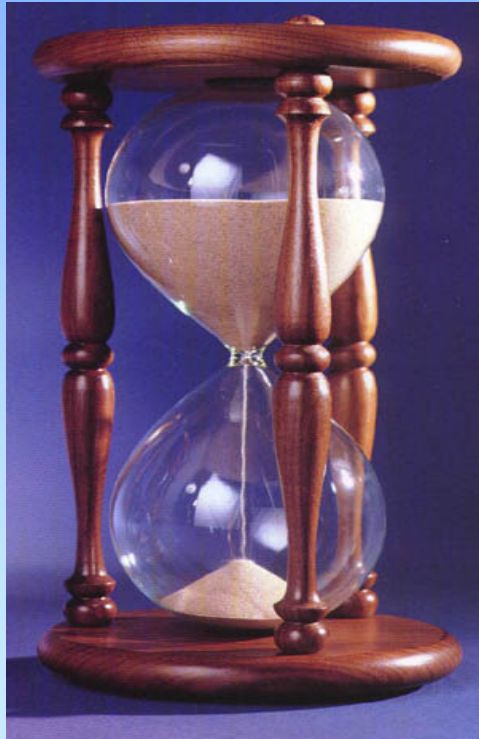
$$f = 0.907$$



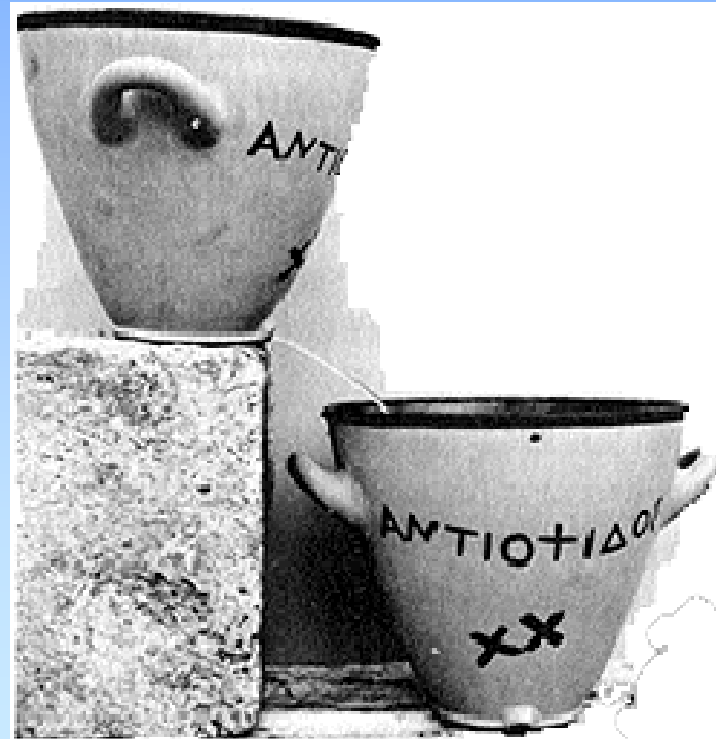
$$f = 0.785$$



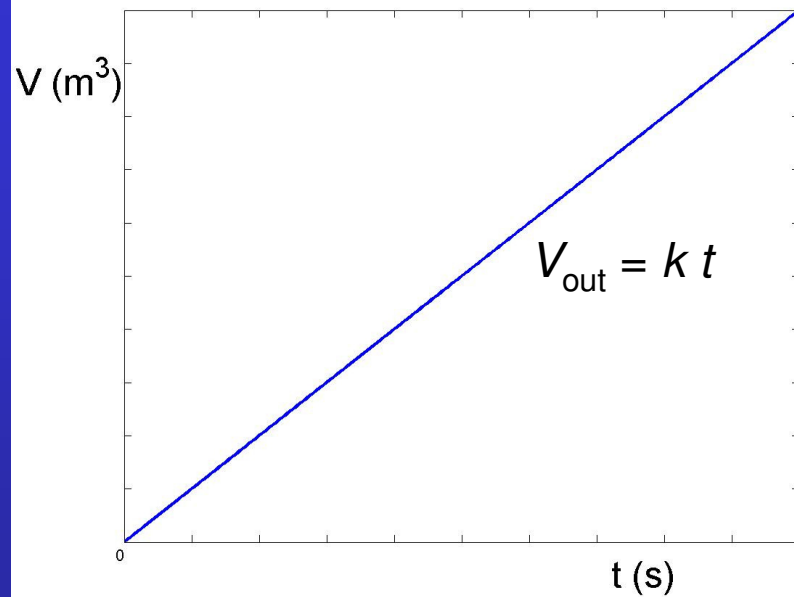
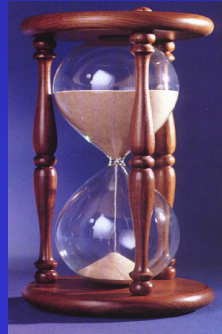
When it behaves like a liquid,  
Granular Matter is a special liquid



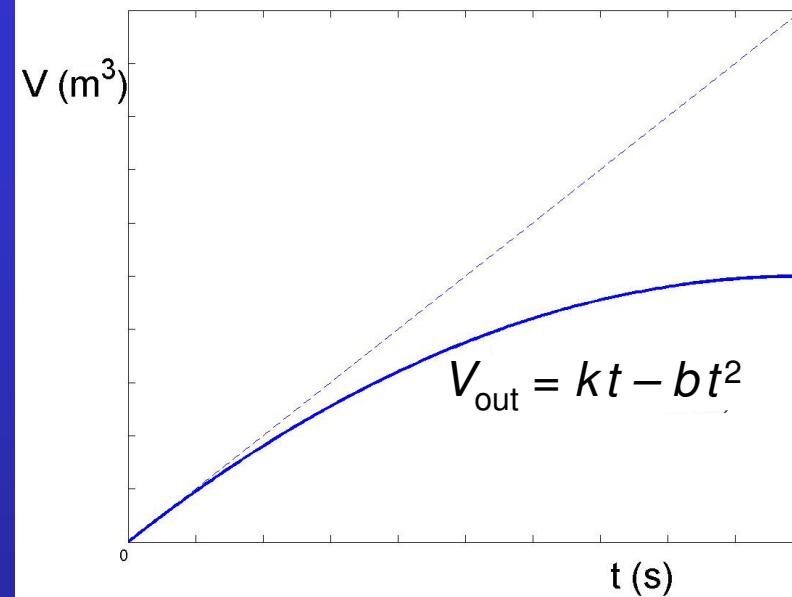
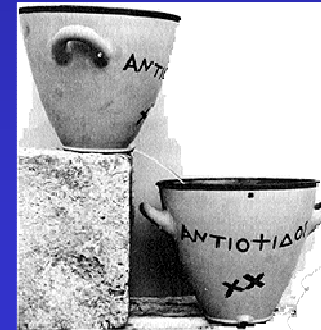
||·?



# hourglass - sand

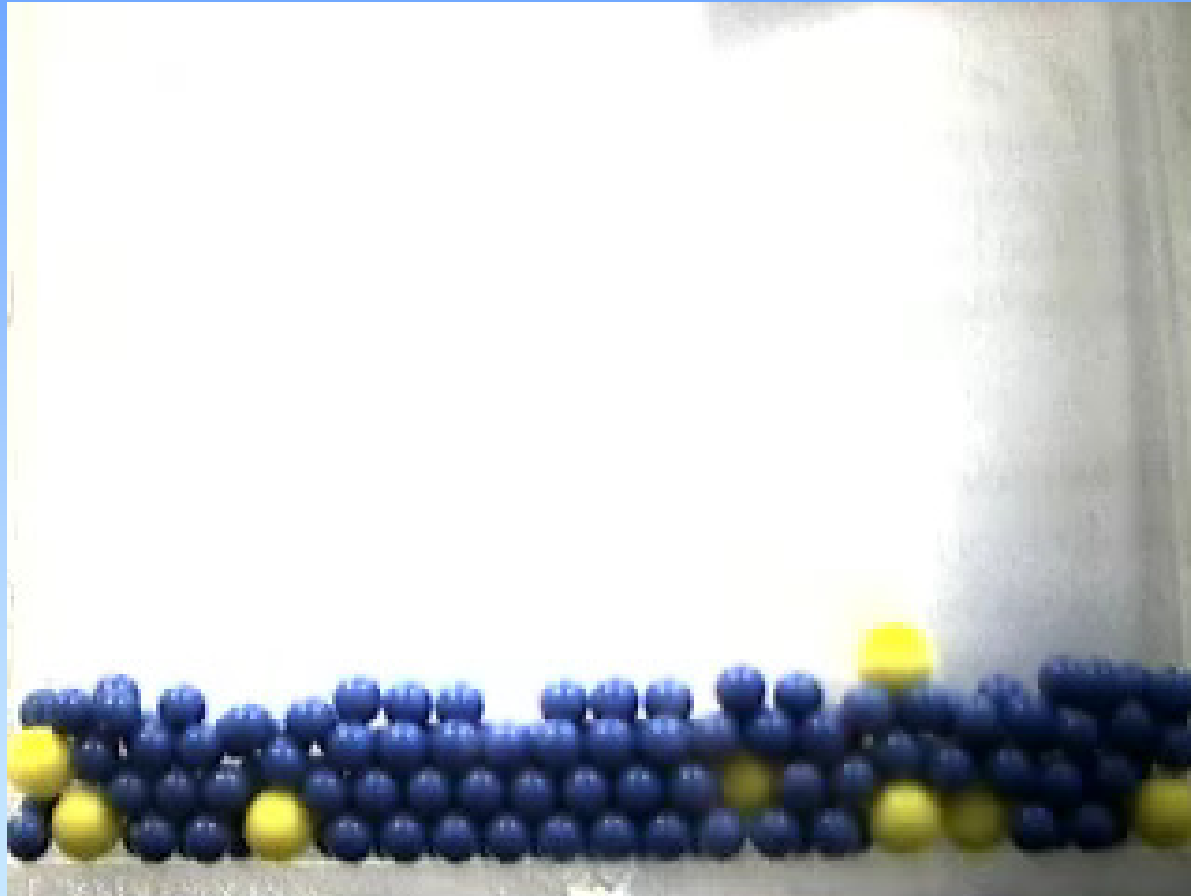


# hourglass - water "klepsydra"





# Vibrated bidisperse mixture



**Segregation !**

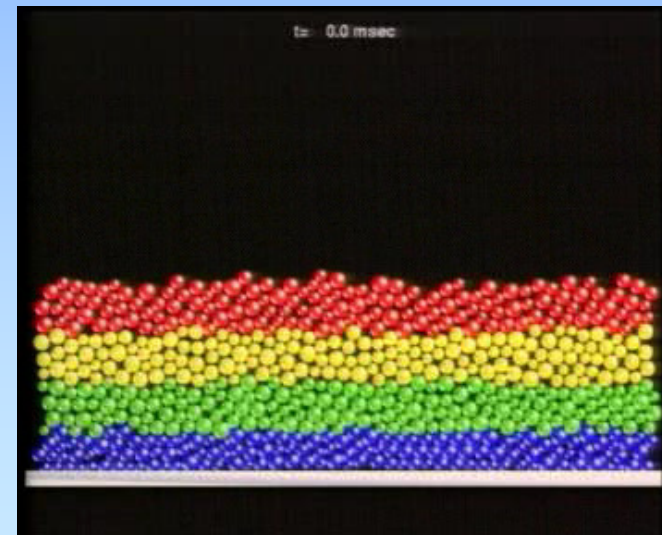
# “Brazil Nut Effect”



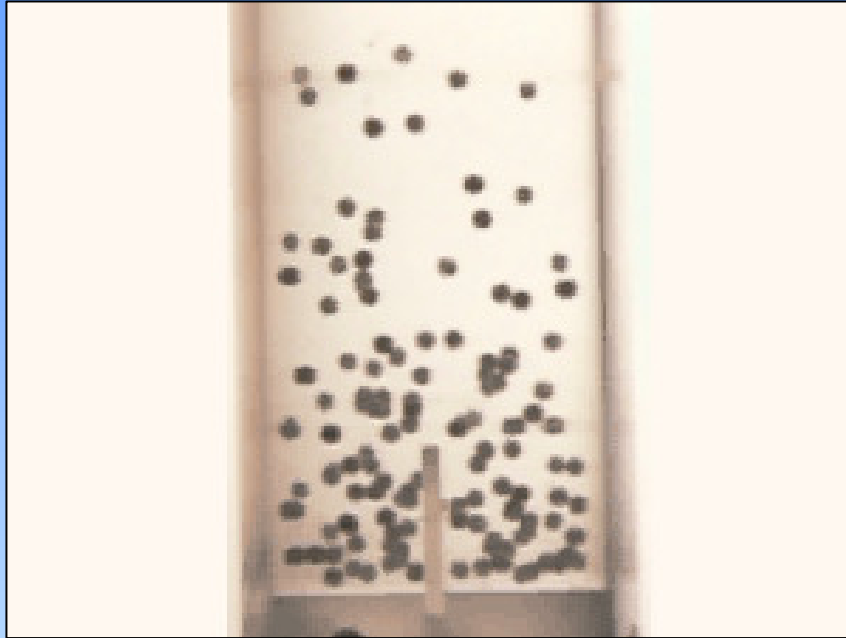


# Three explanations BNE

1. **percolation:** small grains percolate the empty spots between the large ones.
2. **exclusion:** while vibrating small grains fill space below the large ones, not vice versa.
3. **convection:** interaction with walls trigger convection rolls.



large grains can follow the upward,  
but not the downward flow.



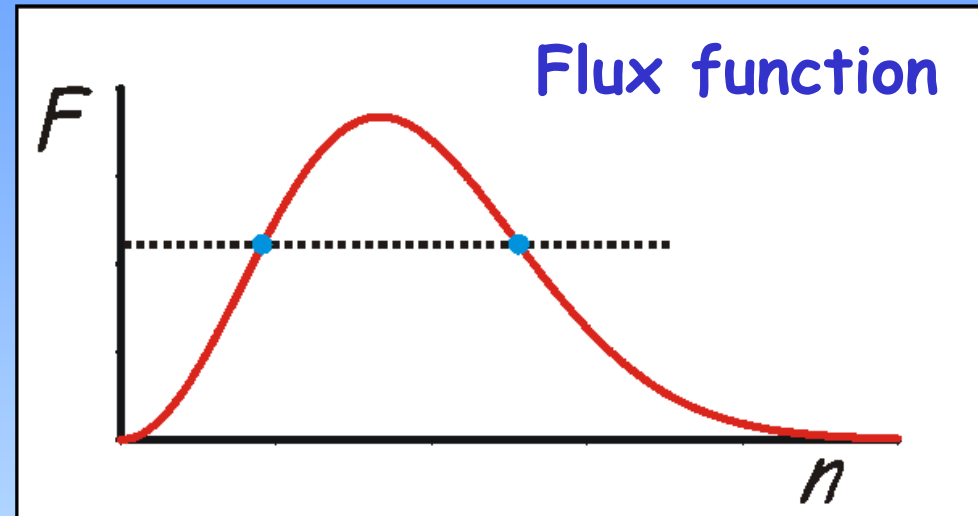
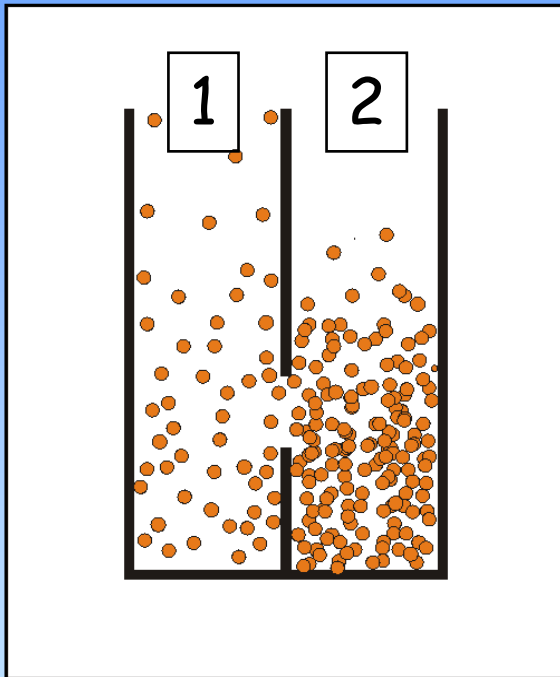
Reason clustering:

Inelastic  
collisions

Driving strength: **high**

And if Granular Matter behaves like a gas,  
it is a special gas

# Flux model



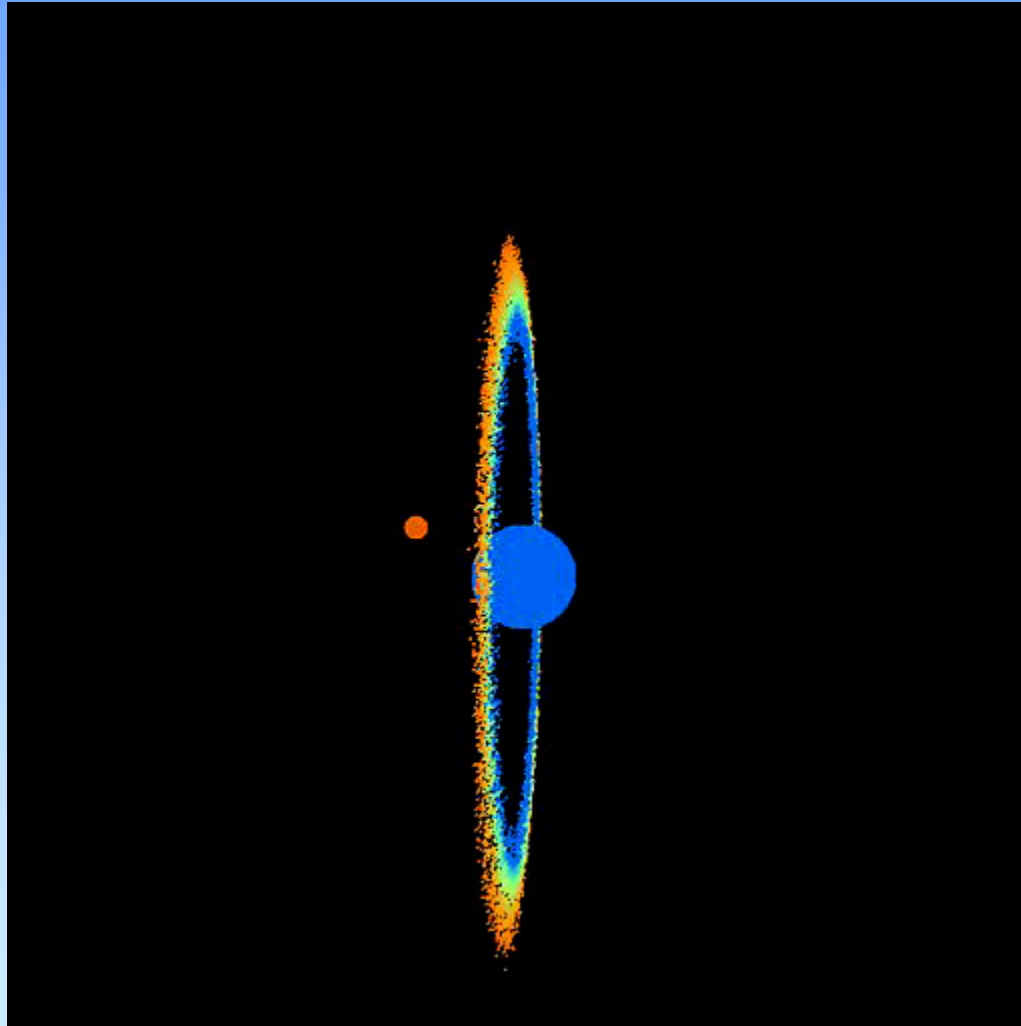
$$\frac{dn_1}{dt} = F_{2 \rightarrow 1} - F_{1 \rightarrow 2}$$

**In five compartments:**





# Planet with rings



# Phenomena

## Granular Solid:

Packing density, dilatancy, force chains, compactification, pressure saturation (RJ law)

## Granular Fluid:

Arching, blocking, convection, segregation

## Granular Gas:

Clustering, non-equipartition

**Why does Granular Matter  
behave so differently  
from other solids and fluids  
we know ?**

# 1. GM is athermal

## *Definition:*

**Granular Matter =**  
many body system in which the typical  
particle size  $> 100 \mu\text{m}$

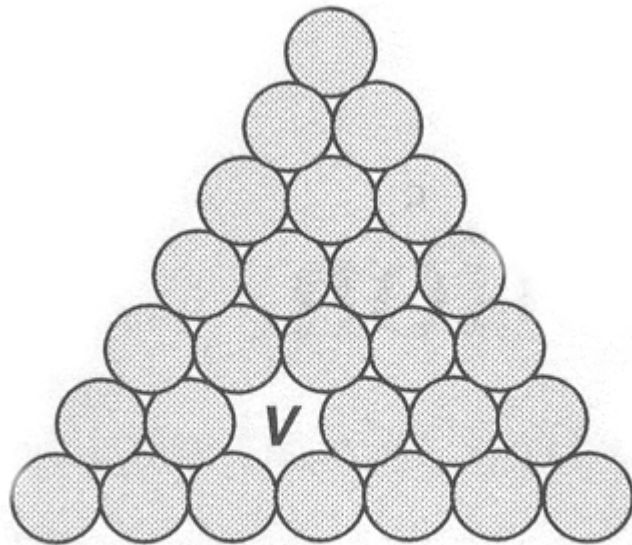
$$\frac{1}{2} m v_{\text{thermal}}^2 = \frac{3}{2} k_B T \quad \longrightarrow \quad (\text{at room temperature})$$

$$v_{\text{thermal}} = \sqrt{\frac{3k_B T}{\frac{4}{3} \pi r^3 \rho}} = \sqrt{\frac{10^{-20}}{10^{-8}}} = 10^{-6} \text{ m/s}$$

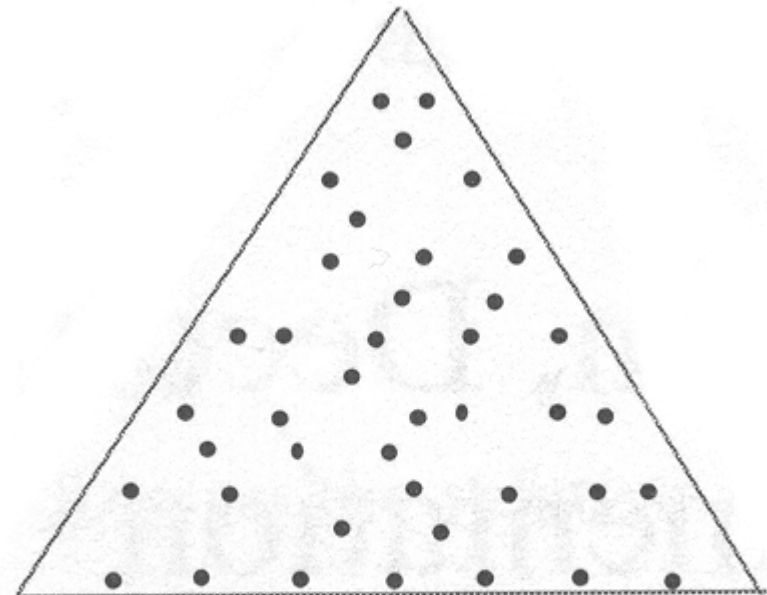
***Thermal energy is negligible for such particles !***



## 2. *GM* interacts through *contact forces*



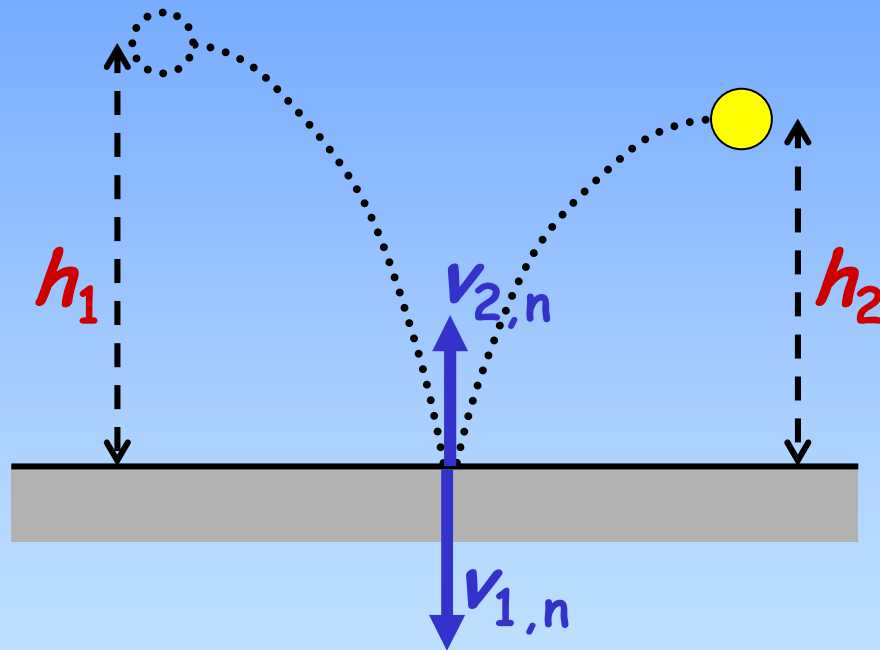
Cannon Ball Stack



Contact Points

*"Chaotic" network of contact points and forces !*

### 3. GM interactions are dissipative



coefficient of  
normal restitution:

$$e = \frac{v_{2,n}}{v_{1,n}} \left( = \sqrt{\frac{h_2}{h_1}} \right)$$

*Grains have many internal degrees of freedom  
through which kinetic energy is dissipated.*

(sound, heat, deformation)

# Implications

1. athermal  $\Rightarrow$  Thermodynamic  $T$  irrelevant

Define granular temperature  $T_g = \left\langle \frac{1}{2} m v_{macro}^2 \right\rangle$

2. contact forces  $\Rightarrow$  Ordered molecular-scale structures do not occur

3. dissipation  $\Rightarrow$  Far-from-equilibrium system

Constant energy supply is necessary to keep systems "alive" (i.e.  $T_g > 0$ )

# Typical practical problems

## Production and handling:

- cornflakes: filling
- pill production: mixing
- casting by sacrificial polystyrene

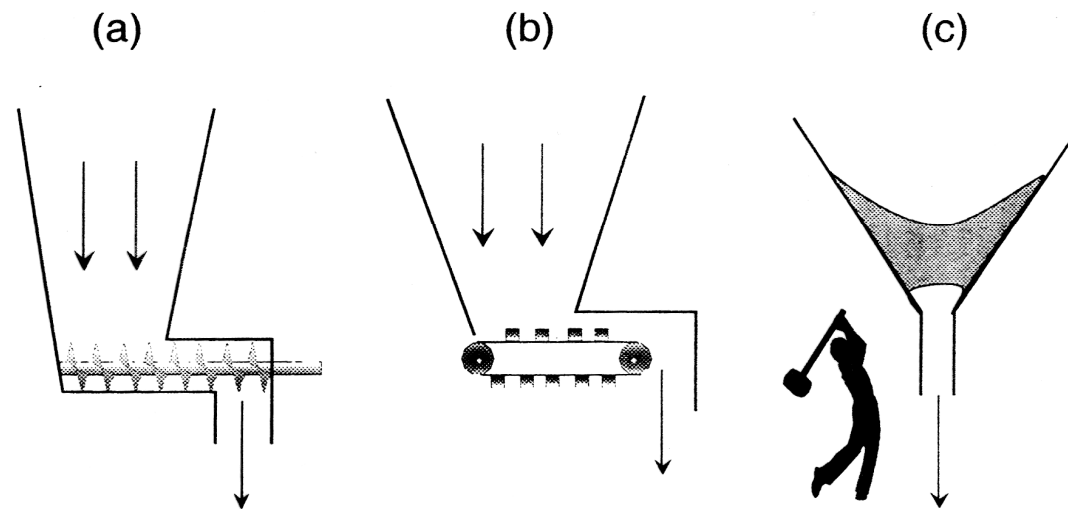
## Nature (geophysics):

- dunes: movement
- avalanches: ranges, volume, prediction
- dikes: stability
- seismology



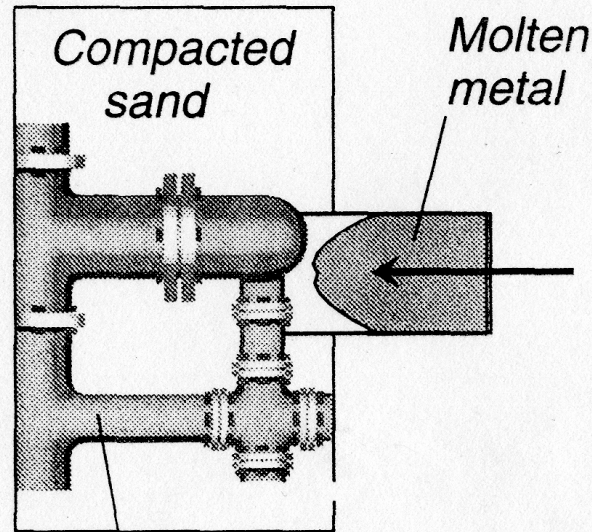
### 1.2.3. *Problems of Segregation*

The industry processes enormous quantities of granular materials year in and year out. Virtually every stage of its operations is at the mercy of segregation—a most irksome phenomenon that tends to separate the components of a mixture supposed

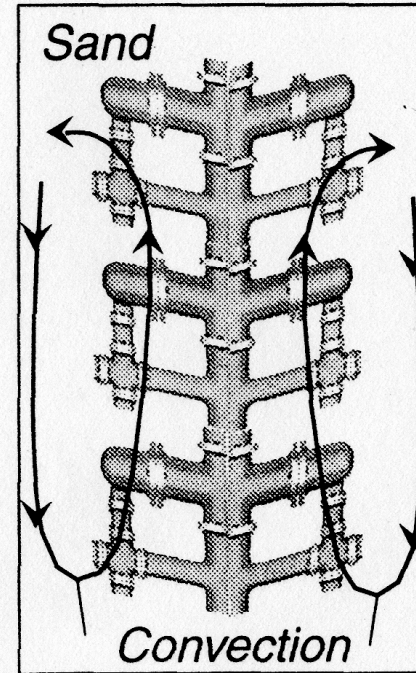


**FIGURE 6.** Three methods used in industry to prevent or remedy blockages by arch effect. They include (a) an Archimedes screw, and (b) a conveyor belt with a corrugated surface. In (c), a plant worker pounds an obstructed hopper with a sledge hammer to get the flow started again. The latter is the method of choice in industries producing low-value-added granular materials.

# Casting by sacrificial polystyrene

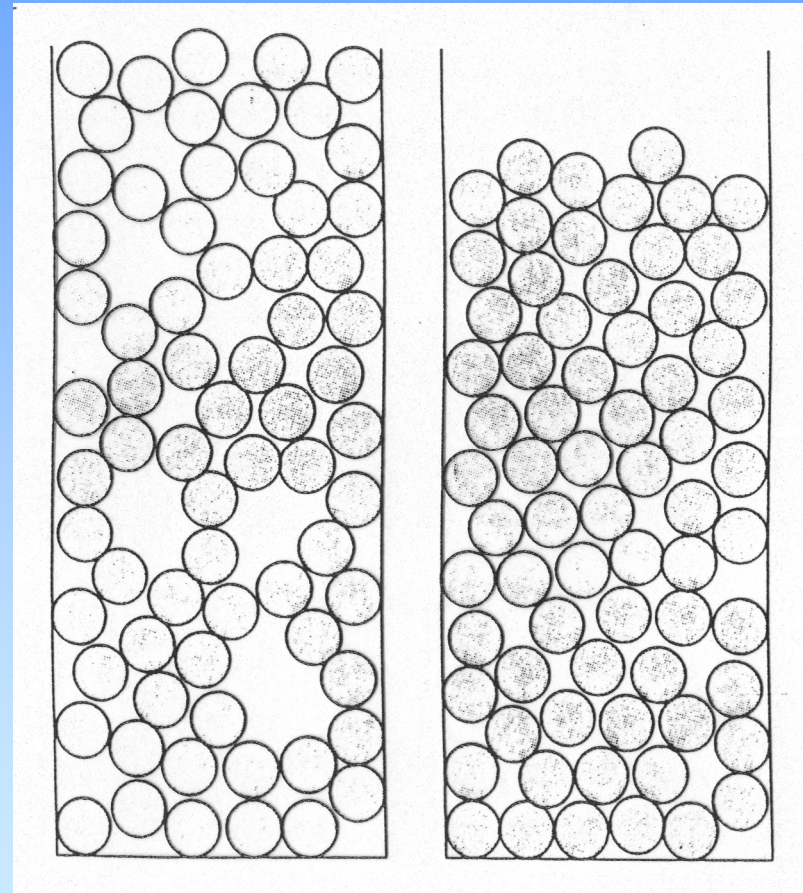
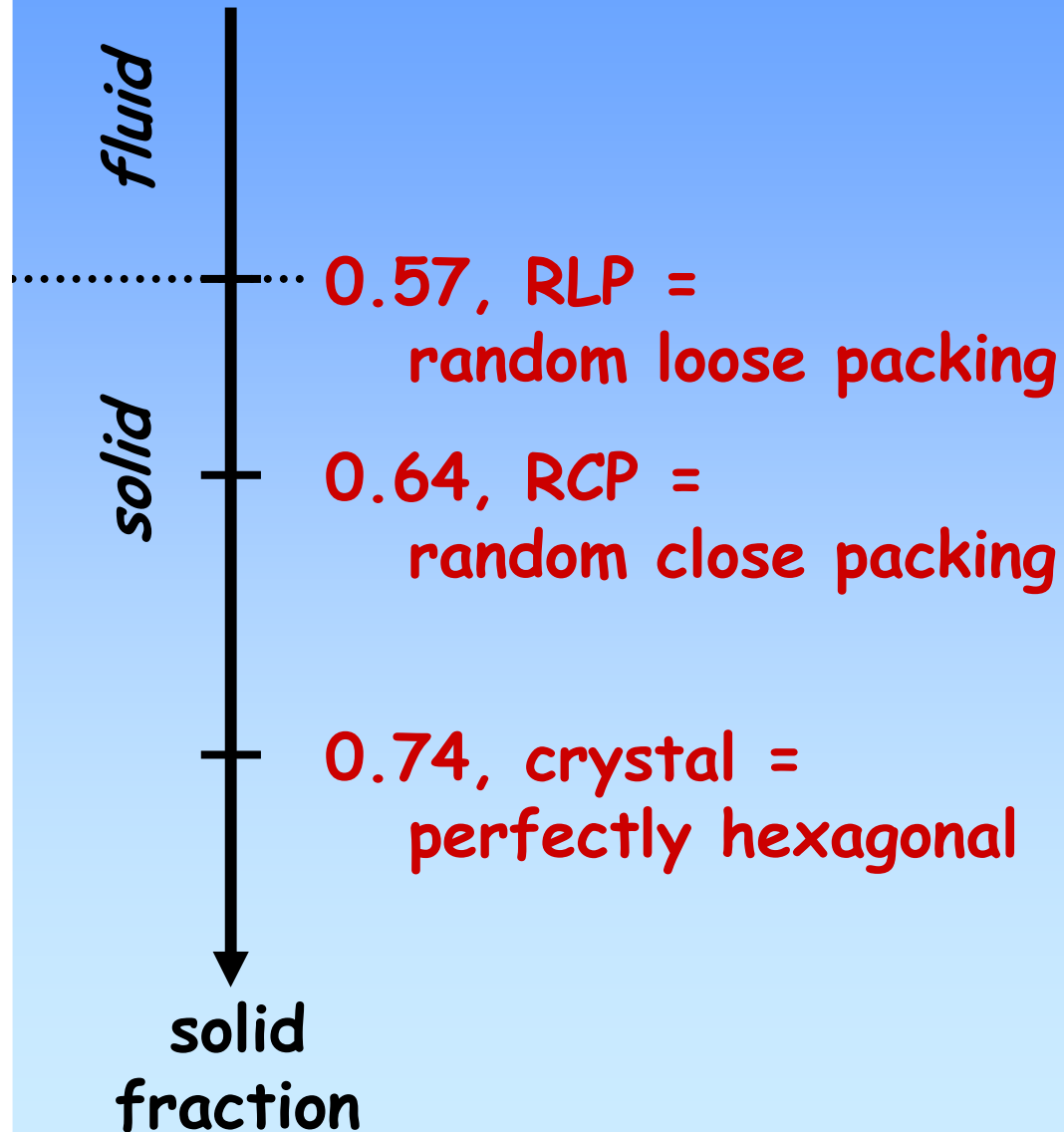


*Polystyrene model*

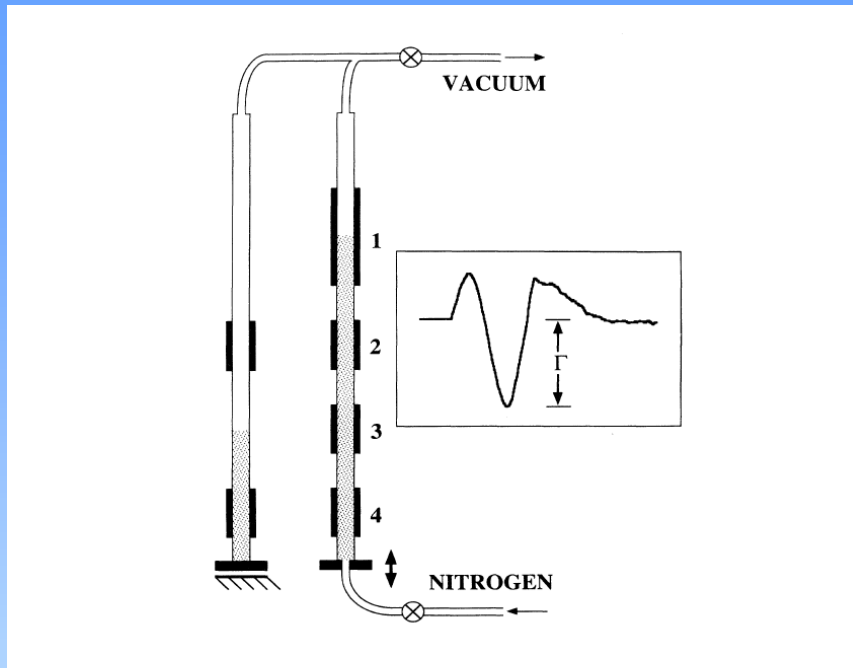


Principle of casting by sacrificial polystyrene.

# Granular packing (for spheres)



# Compactification experiment

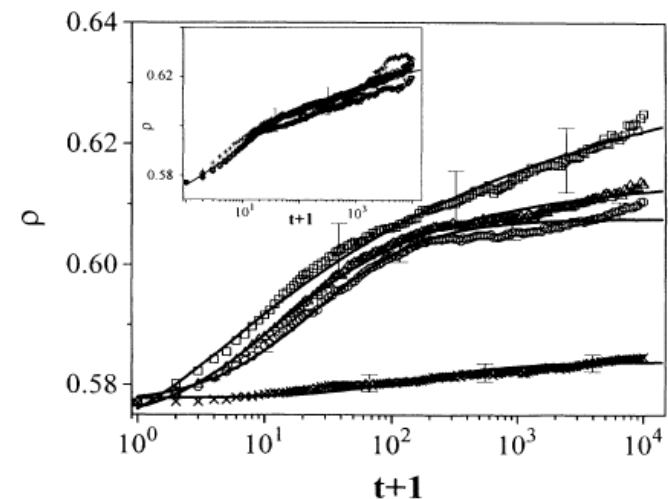
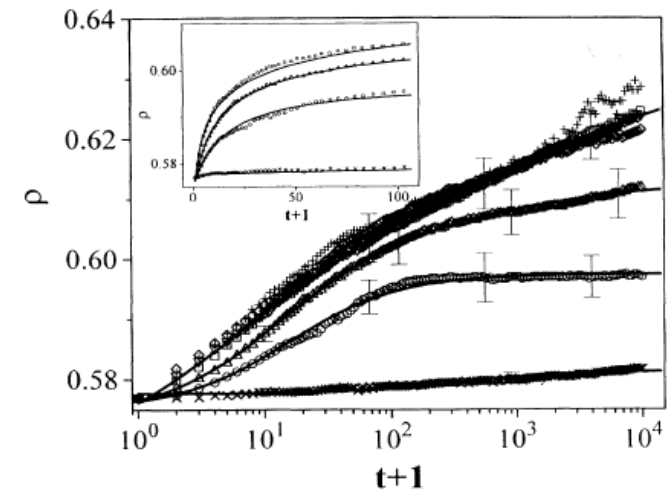


$$\rho(t) = \rho_f - \frac{\rho_f - \rho_0}{1 + B \log(1 + t / \tau)}$$

$$\rho(t = 0) = \rho_0; \quad \rho(t \rightarrow \infty) = \rho_f$$

regime 1: *local reorganization*

regime 2: *global reorganization*





# Analogy: car-parking in street

## Model (Ben-Naim):

- Initial state: randomly parked cars (no extra fit in)
- Start to move cars randomly. Whenever there is a large enough gap, a new car jumps in.

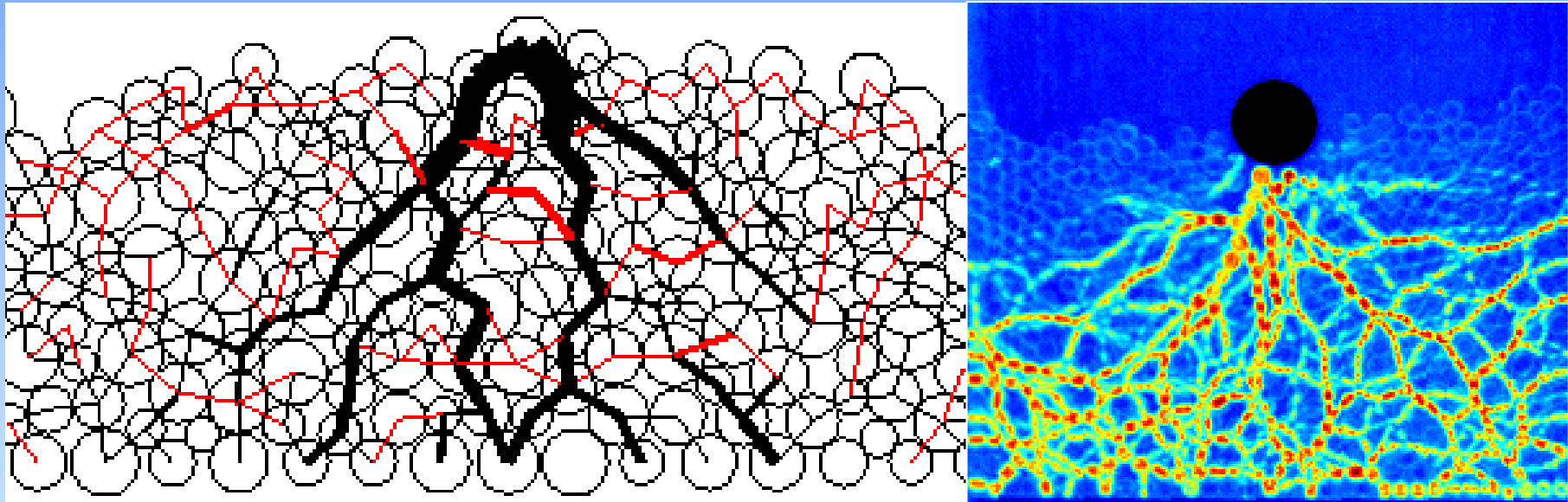
regime 1: movement of a single car creates gap

regime 2: more than one car has to move:

required time for gap to open grows exponentially:

$$\frac{\rho(t) - \rho_f}{\rho_f - \rho_0} \propto \frac{1}{\log(t / \tau)}$$

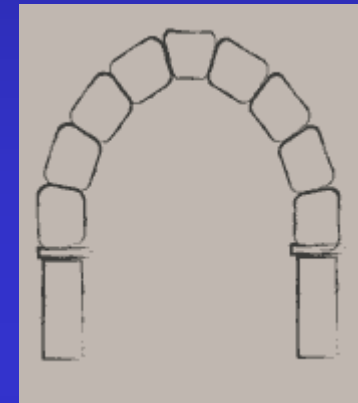
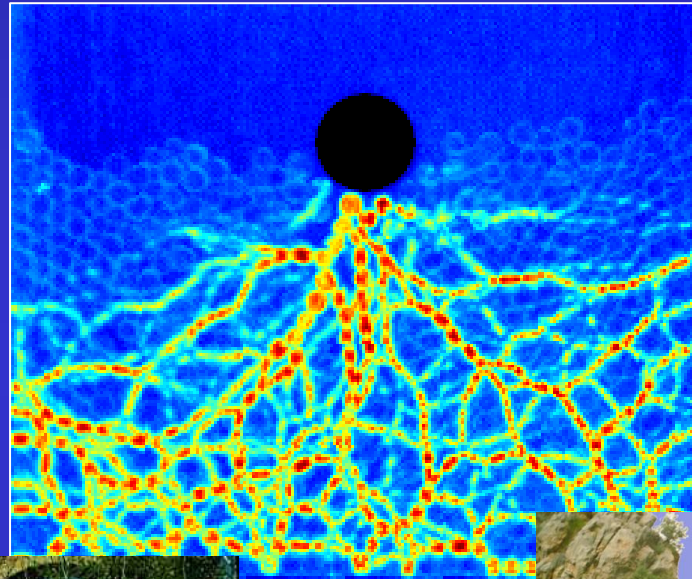
# Force Chains



(Bob Behringer, Duke)

# In stalling flow, force chains manifest themselves as arches

Segovia, Spain



Pont du Gard, France



Spain

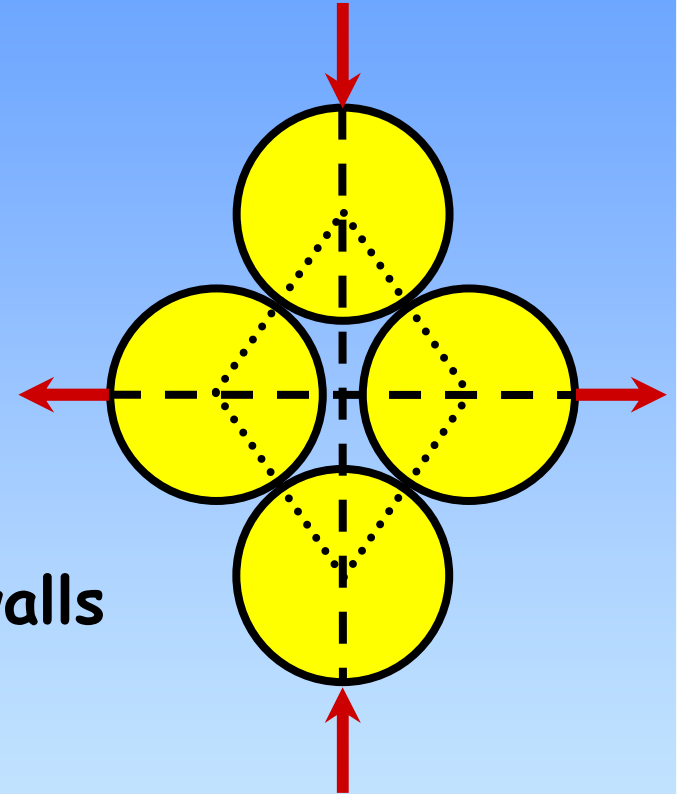
# Importance of sidewalls: Rayleigh-Janssen model

Force parallelogram as unit cell  
of a 2D granular medium



vertical forces  $\Rightarrow$  horizontal forces

balanced by sidewalls



*Lord Rayleigh:*

$$p_h = K p_v$$

$K$  = coefficient of *redirection*

# Importance of sidewalls: Rayleigh-Janssen model (2)

Slice experiences friction  
force with sidewalls:

$$\begin{aligned}dF_{\text{friction}} &= \mu_s p_h U dh \\ &= \mu_s K p_v U dh\end{aligned}$$

Vertical force balance on slice:

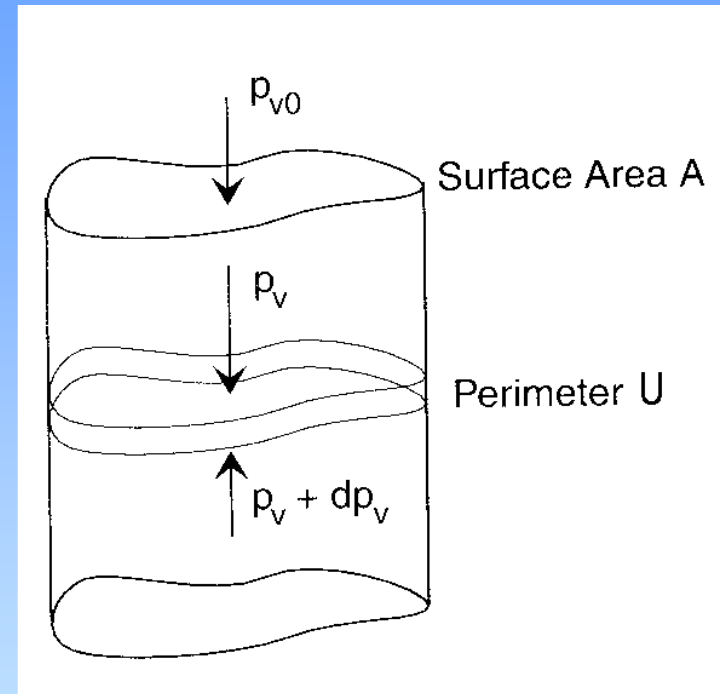
$$\begin{aligned}(p_v(h + dh) - p_v(h))A + \\ \mu_s K p_v U dh = \rho g A dh\end{aligned}$$

$$\frac{dp_v}{dh} + \mu_s K \frac{U}{A} p_v = \rho g$$

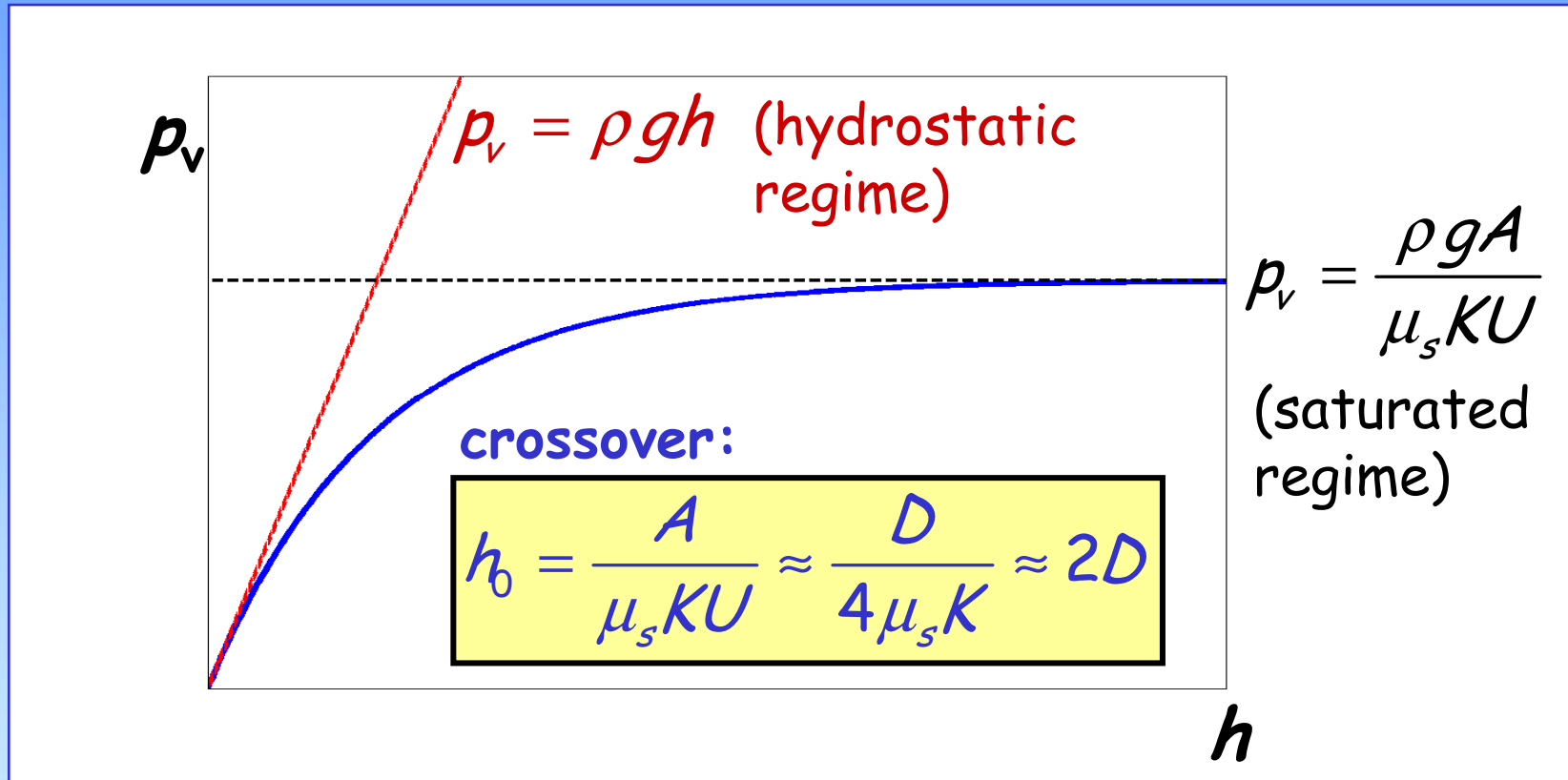
Integration gives:

$$p_v = \frac{\rho g A}{\mu_s K U} \left( 1 - \exp\left(-\mu_s K \frac{U}{A} h\right) \right)$$

Janssen's  
equation



# Importance of sidewalls: Rayleigh-Janssen model (2)



$$p_v = \frac{\rho g A}{\mu_s K U} \left( 1 - \exp\left(-\mu_s K \frac{U}{A} h\right) \right)$$

Janssen's  
equation



# Effective weight of granulate in silo

$$\chi = \frac{\mu_s KU}{A} h \quad (\text{decompaction parameter})$$

Effective weight on bottom =  $F_v(h) = p_v(h) A$

$$F_v(h) = mg \frac{1 - \exp(-\chi)}{\chi} \approx \frac{mg}{\chi} \quad (\text{for large } \chi, \text{ i.e., large } h)$$

What happens to the remaining weight?

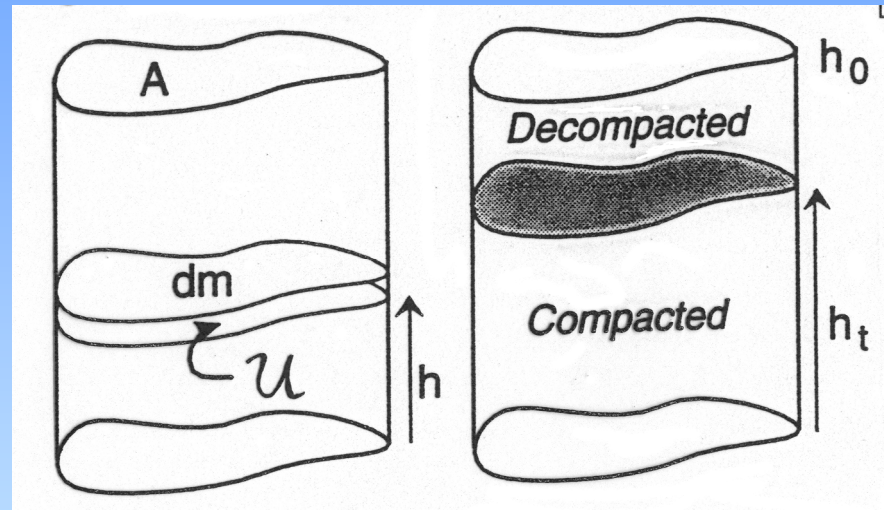
# Collapsing silos

Walls take this weight!



# Decompactification through shaking

Shaking:  $a \sin(\omega t)$ ; dim.less acceleration:  $\Gamma = \frac{a\omega^2}{g}$



“Decompacted” means: acceleration overcomes friction

Force balance: acceleration - gravity = (wall) friction:

$$\Gamma g dm - g dm = dF_{\text{friction}}$$

*Sand moves freely if lhs > rhs !*

# Decompactification through shaking (threshold calculation)

Total height of stack:  $h_0$

Threshold condition lhs > rhs fulfilled from  $h_t$  ( $< h_0$ ) on.

$$(\Gamma - 1) = \frac{1}{g} \frac{dF_{\text{friction}}}{dm} = \frac{\mu_s K p_v(h_t) U dh}{g \rho A dh} = \frac{\mu_s K U}{\rho g A} p_v(h_t)$$

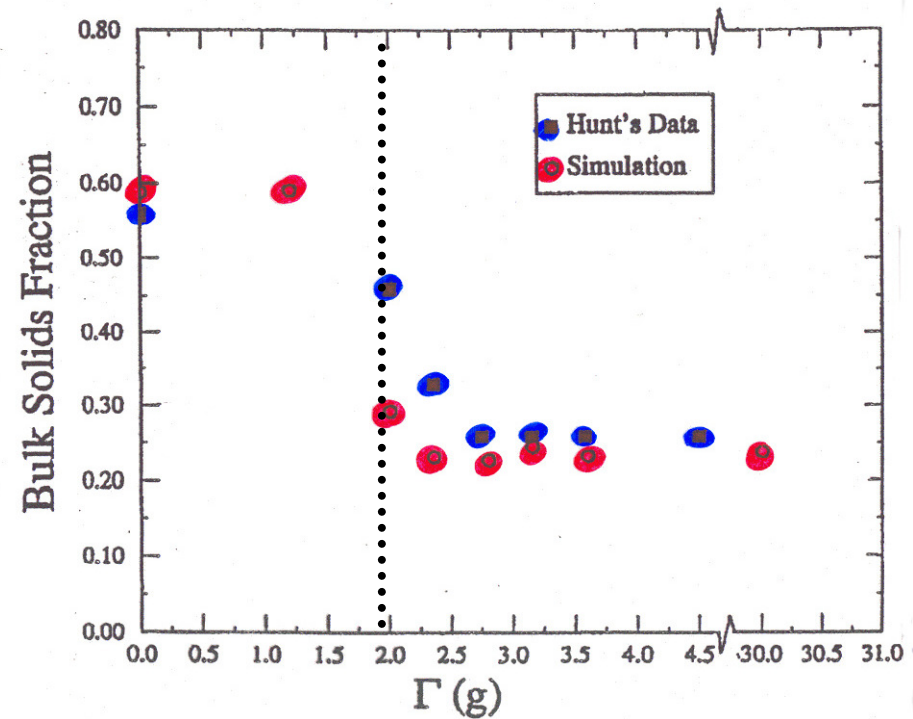
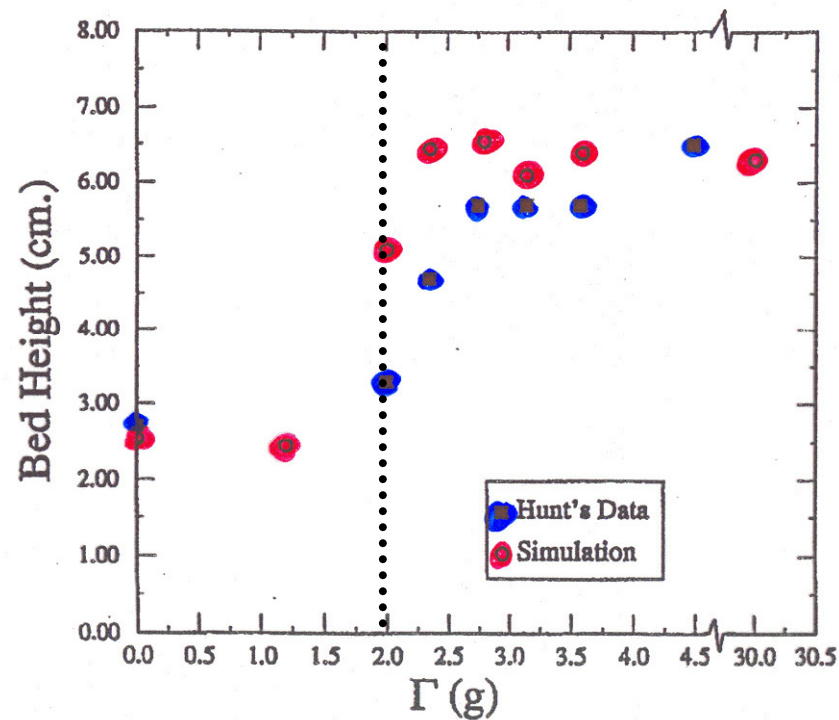
$$p_v(h_t) = \frac{\rho g A}{\mu_s K U} \left( 1 - \exp\left(-\mu_s K \frac{U}{A} (h_0 - h_t)\right) \right)$$

$$\Gamma = 2 - \exp\left(-\frac{K \mu_s U}{A} (h_0 - h_t)\right)$$

$\Gamma = 1$  means:  $h_t = h_0$ ; nothing can be fluidized

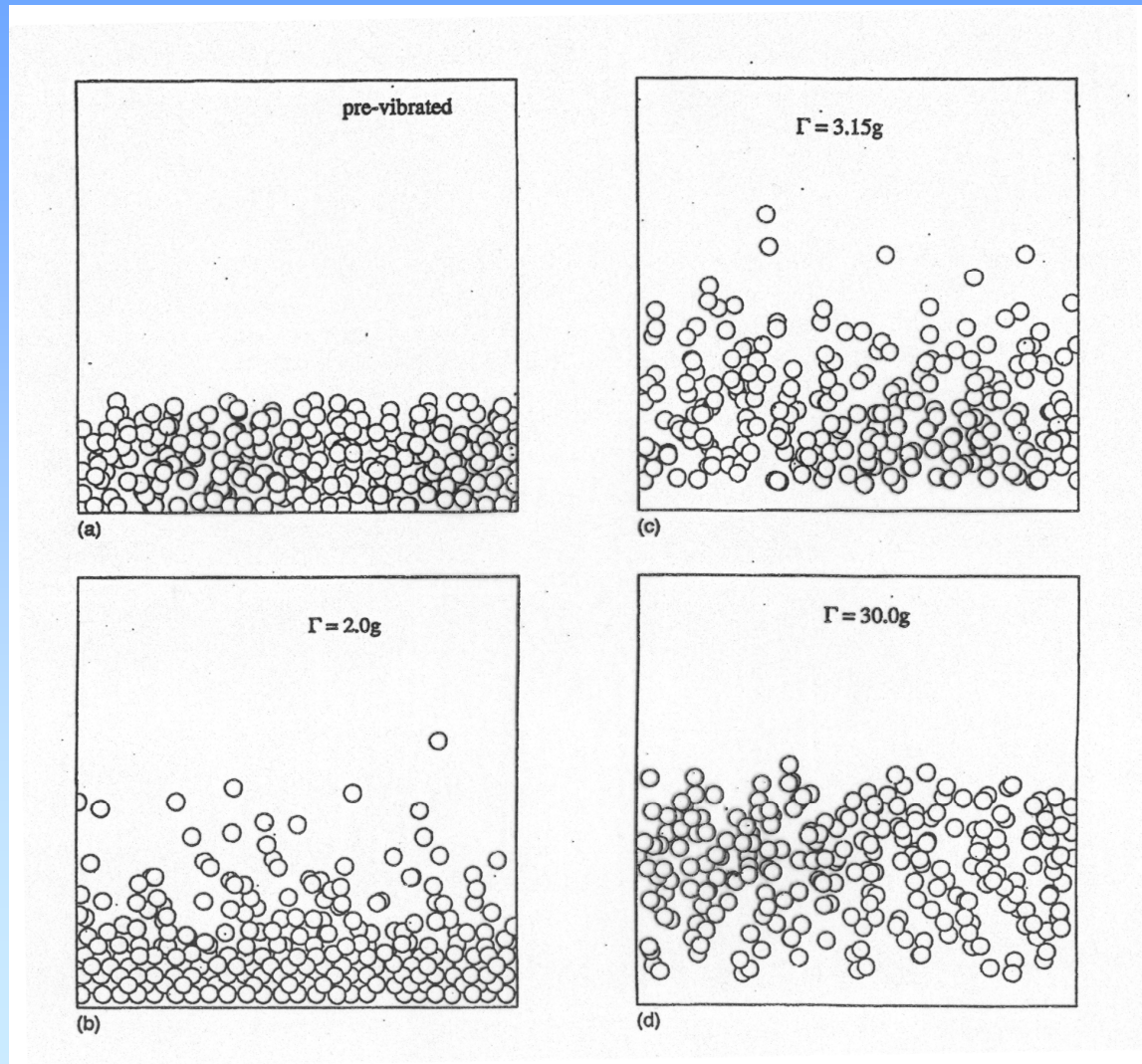
$\Gamma = 2$  or larger: all can be fluidized

# Decompaction: Experiment, Simulation and Theory



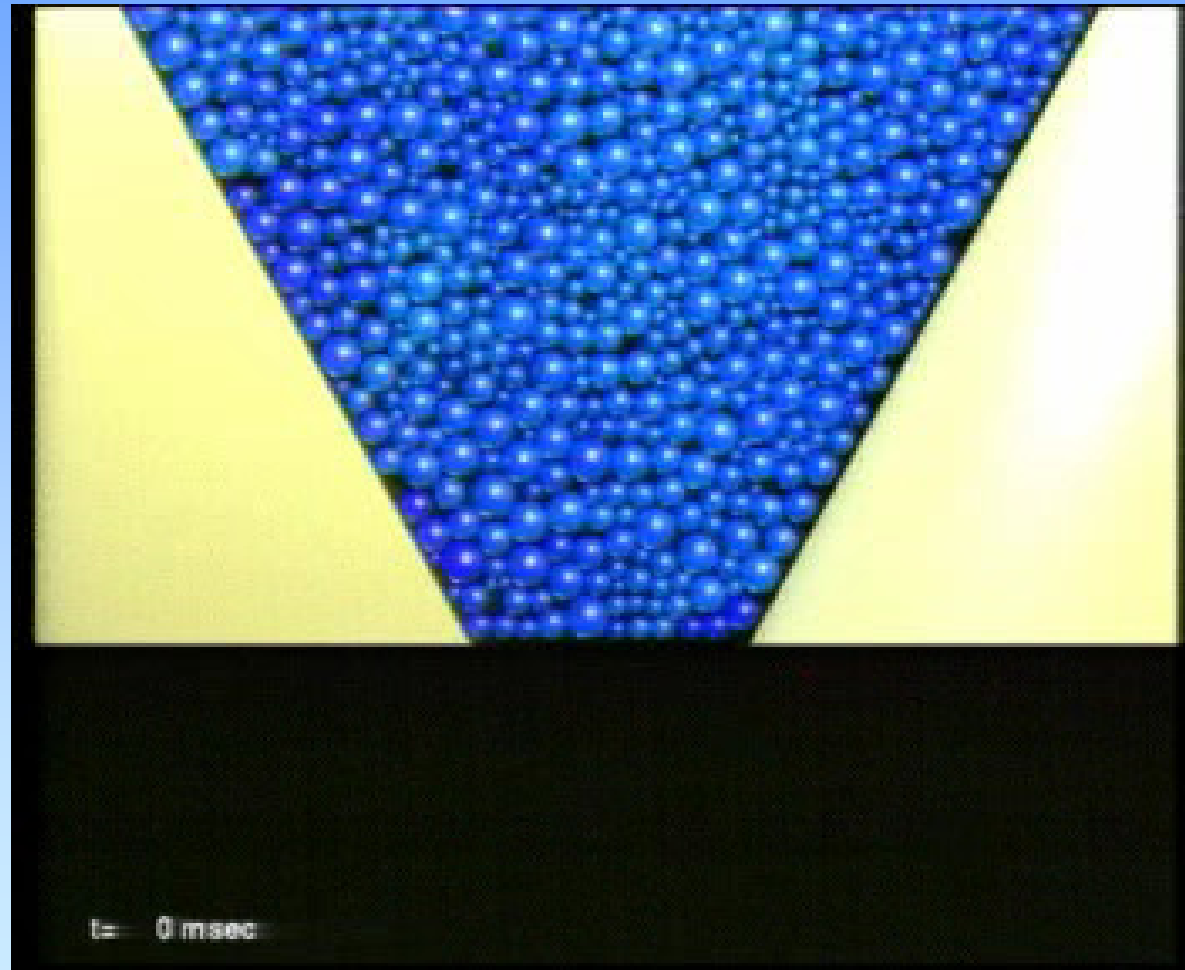
Lan & Rosato, Phys. Fluids 7, 1818 (1995)

# Decompaction: Experiment, Simulation and Theory

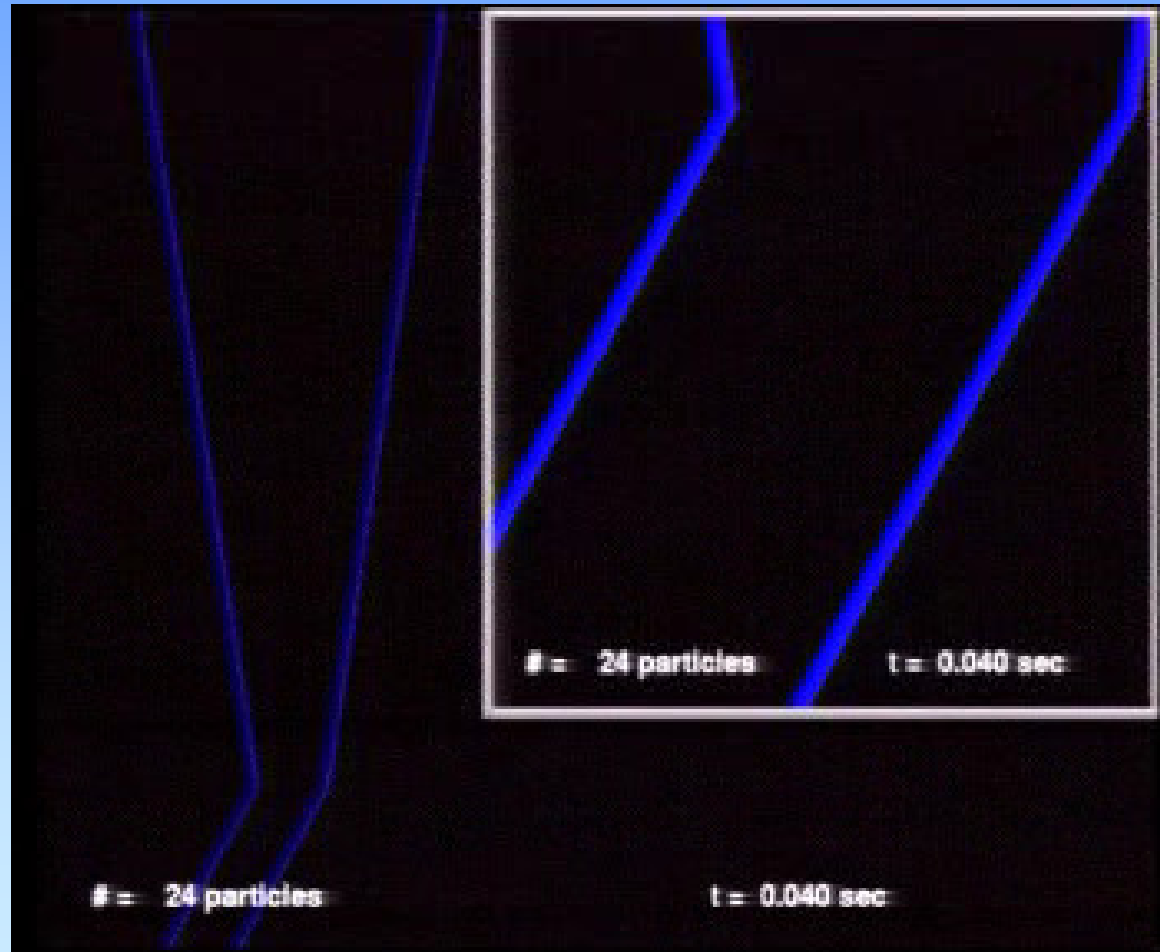




# Granular matter in a hopper...



... and in a funnel



# Is a general hydrodynamic description of granular matter possible?

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PHYSICAL REVIEW LETTERS

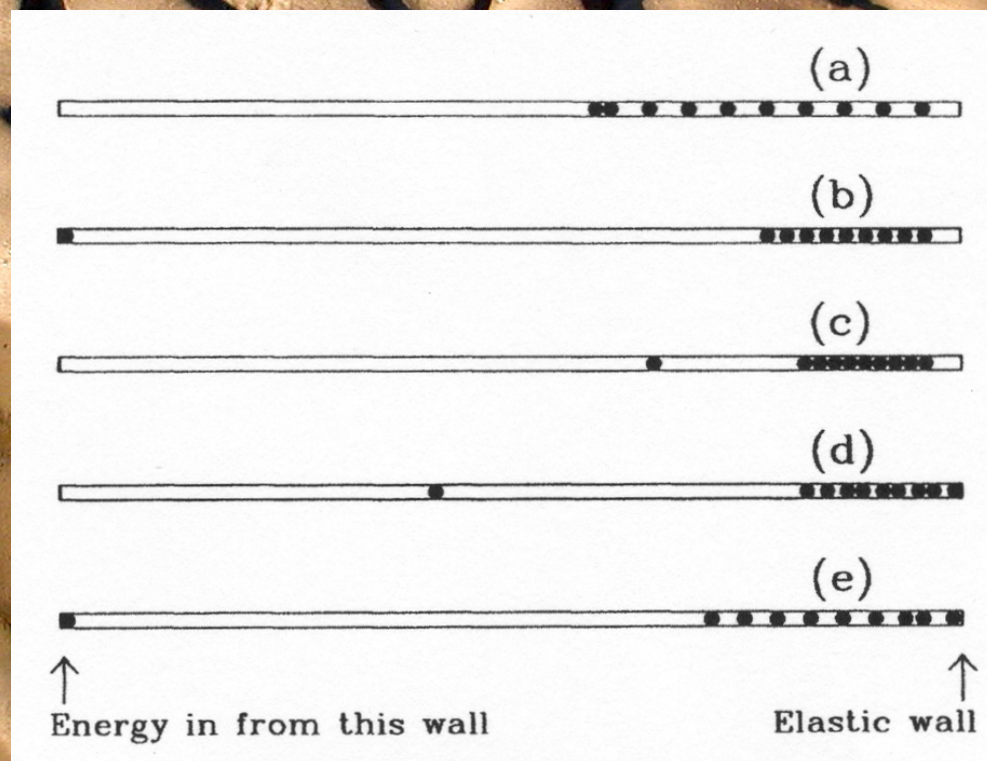
20 FEBRUARY 1995

## Breakdown of Hydrodynamics in a One-Dimensional System of Inelastic Particles

Yunson Du, Hao Li, and Leo P. Kadanoff

*The James Franck Institute, The University of Chicago, Chicago, Illinois 60637*

(Received 15 August 1994)



# A) Hydrodynamic approach

Coarse graining over small intervals  $\Delta x$ ,  $\Delta t$  to define macroscopic quantities:

**density:**

$$\rho(x, t) = \left\langle \sum_i \delta(x_i(t) - x) \right\rangle_{\Delta x, \Delta t}$$

**velocity:**

$$U(x, t) = \left\langle \sum_i v_i(t) \delta(x_i(t) - x) \right\rangle_{\Delta x, \Delta t}$$

**temperature:**

$$T(x, t) = \left\langle \sum_i (v_i(t) - U(x, t))^2 \delta(x_i(t) - x) \right\rangle_{\Delta x, \Delta t}$$

Assuming local “thermal” equilibrium, one can derive mass, momentum, and energy conservation laws:



# Conservation laws

In the dilute limit, using the ideal gas law:

$$\cancel{\partial_t \rho} + \cancel{\partial_x (\rho u)} = 0$$

$$\cancel{\rho \partial_t u} + \cancel{\rho u \partial_x u} = -c_1 \partial_x (\rho T)$$

$$\cancel{\rho \partial_t T} + \cancel{\rho u \partial_x T} + \cancel{c_1 \rho T \partial_x u} - c_2 \partial_x^2 (T^{3/2}) = -c_3 \varepsilon \rho^2 T^{3/2}$$

$\varepsilon = (1 - e) / 2$  expresses *inelasticity*

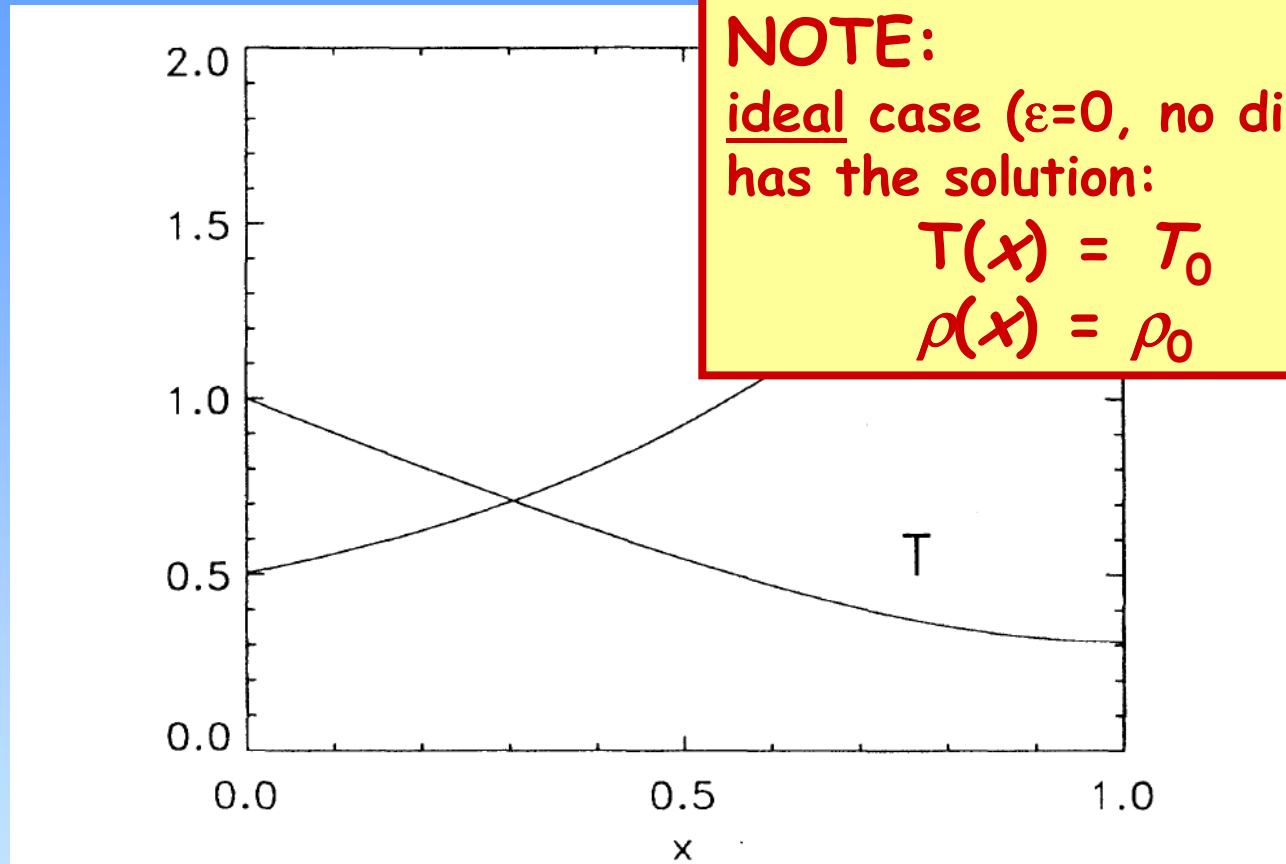
In the stationary limit ( $u=0$ ,  $d_t=0$ ) this becomes:

$$\rho T = \text{constant}$$

$$\partial_x^2 (T^{3/2}) = \frac{c_3 \varepsilon}{c_2} \rho^2 T^{3/2}$$

These equations can be solved analytically:

# Hydrodynamic solution:



**NOTE:**

ideal case ( $\varepsilon=0$ , no dissipation)  
has the solution:

$$T(x) = T_0$$

$$\rho(x) = \rho_0$$

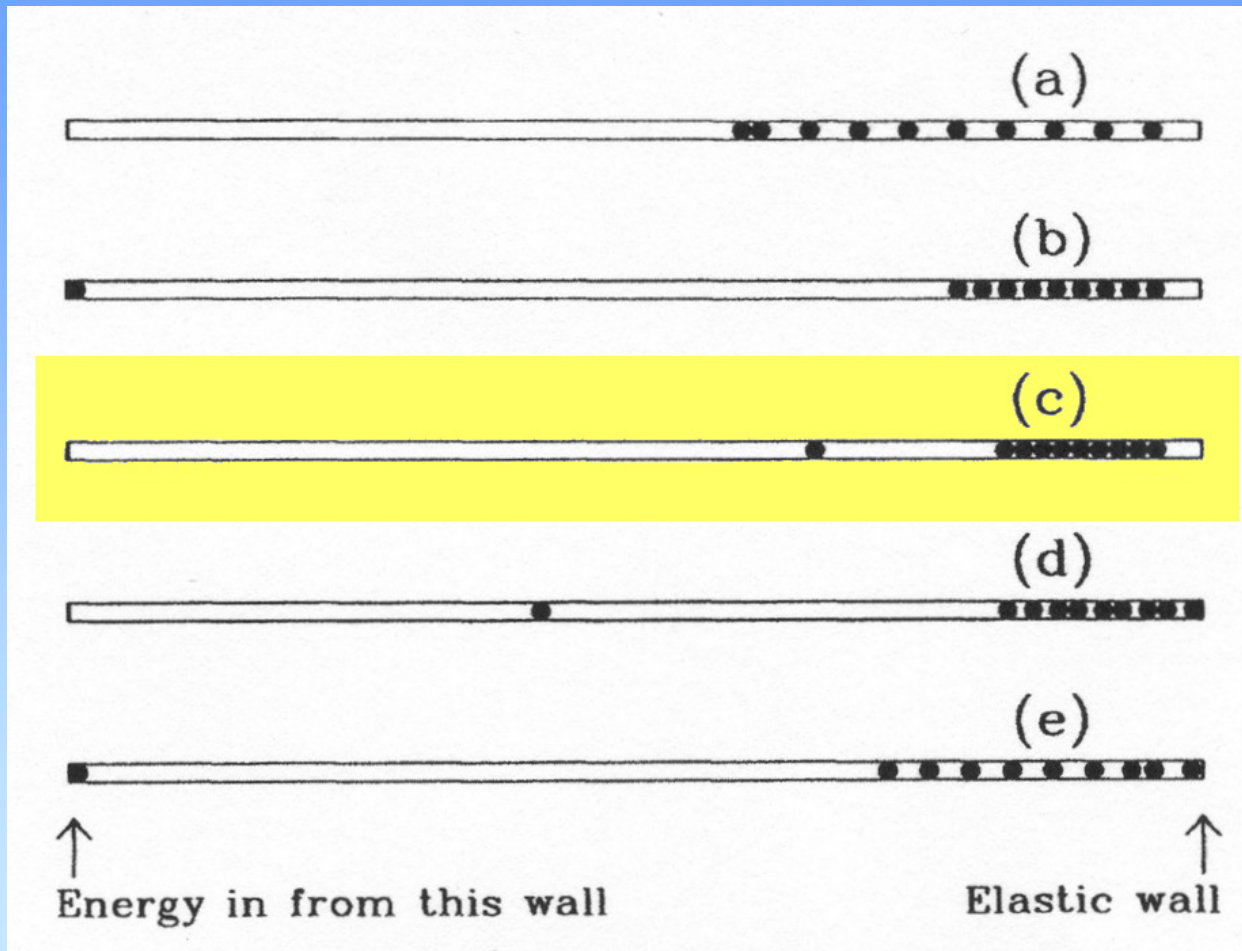
Using the boundary conditions:

$$T(0) = T_0 \quad \text{[constant } T \text{ at left border]}$$

$$\partial_x T(1) = 0 \quad \text{[elastic wall (no heat flux) at right border]}$$



# Particle dynamics solution:



(using MD simulations)

# B) Discrete description

2-particle collision with

\* **momentum conservation:**

$$v_1 + v_2 = v_1' + v_2'$$

\* **energy dissipation:**

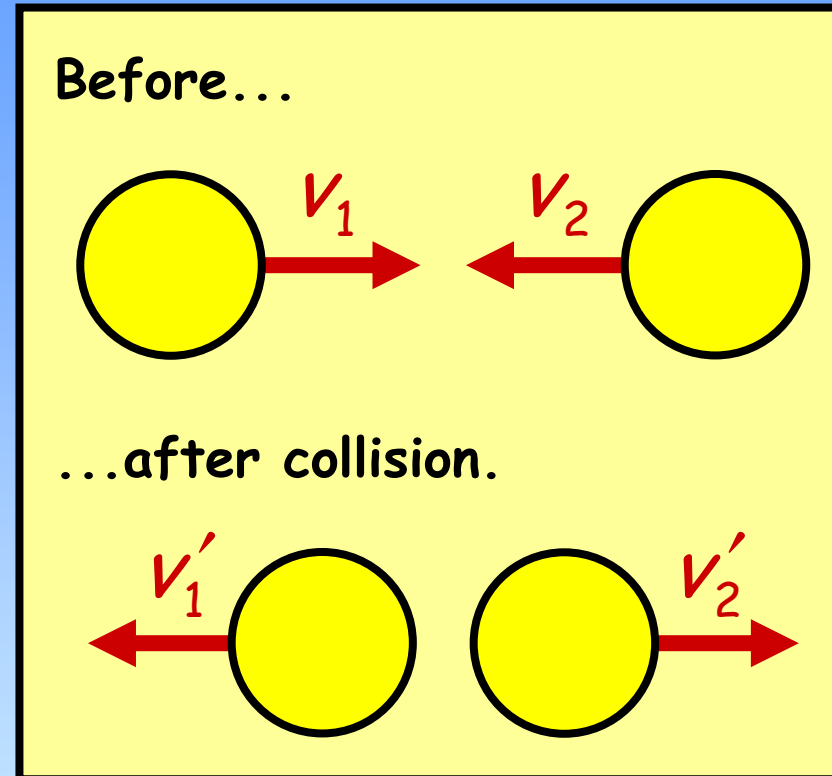
$$v_1' - v_2' = -e(v_1 - v_2)$$

This implies:

$$v_1' = \varepsilon v_1 + (1 - \varepsilon)v_2$$

$$v_2' = (1 - \varepsilon)v_1 + \varepsilon v_2$$

with:  $\varepsilon = (1 - e)/2$



**Ideal case  $\varepsilon = 0$ :**

$V_1' = V_2$ ;  $V_2' = V_1$ , exchange of velocities.

Finally all velocities will be given by the PDF of velocities on the left.

Uniform distribution of particles, consistent with continuum description.

**Non-ideal case  $\varepsilon > 0$ :**

**Numerical result very different from continuum result!**

1 fast particle  $v_N \sim \sqrt{T_0}$  and  $(N-1)$  slow particles, clustering to the right and dissipating energy. Fast particle transports energy from left to right.

No longer local "thermal" equilibrium !

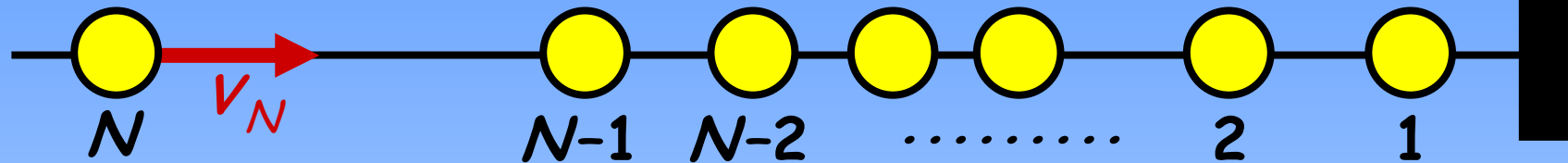
**Breakdown of continuum approach !**

# Velocity center of

$$v_1' = \varepsilon v_1 + (1 - \varepsilon)v_2$$

$$v_2' = (1 - \varepsilon)v_1 + \varepsilon v_2$$

assume:  $v_0 = \text{const} = 1$  (no random distribution)



before first collision:

$$v_N = v_0 = 1, v_i = 0 \text{ for } i < N$$

after first collision (between  $N$  and  $N-1$ ):

$$v_N = \varepsilon, v_{N-1} = 1 - \varepsilon, v_i = 0 \text{ for } i < N-1$$

after second collision (between  $N-1$  and  $N-2$ ):

$$v_N = \varepsilon, v_{N-1} = (1 - \varepsilon)\varepsilon, v_{N-2} = (1 - \varepsilon)^2, v_i = 0 \text{ for } i < N-2$$

...

after  $(N-1)^{\text{th}}$  collision (between 2 and 1):

$$v_N = \varepsilon, v_{N-1} = (1 - \varepsilon)\varepsilon, v_{N-2} = (1 - \varepsilon)^2 \varepsilon, \dots,$$

$$v_2 = (1 - \varepsilon)^{N-2} \varepsilon, v_1 = (1 - \varepsilon)^{N-1}$$

# Velocity center of mass (2)

Mean velocity of particles  $N, N-1, \dots, 3, 2$ :

$$\begin{aligned}V_{\text{CM-cluster}} &= V_N + V_{N-1} + V_{N-2} + \dots + V_3 + V_2 \\&= \varepsilon + (1 - \varepsilon)\varepsilon + (1 - \varepsilon)^2\varepsilon + \dots + (1 - \varepsilon)^{N-2}\varepsilon \\&= \varepsilon \sum_{k=0}^{N-2} (1 - \varepsilon)^k = 1 - (1 - \varepsilon)^{N-1} \\&\approx 1 - \exp(-(N - 1)\varepsilon)\end{aligned}$$

for large  $N$

\*  $\varepsilon = 0$ , ideal case:

$$V_{\text{CM-cluster}} = 0$$

\*  $\varepsilon \neq 0$ , real case:

$$V_{\text{CM-cluster}} > 0$$

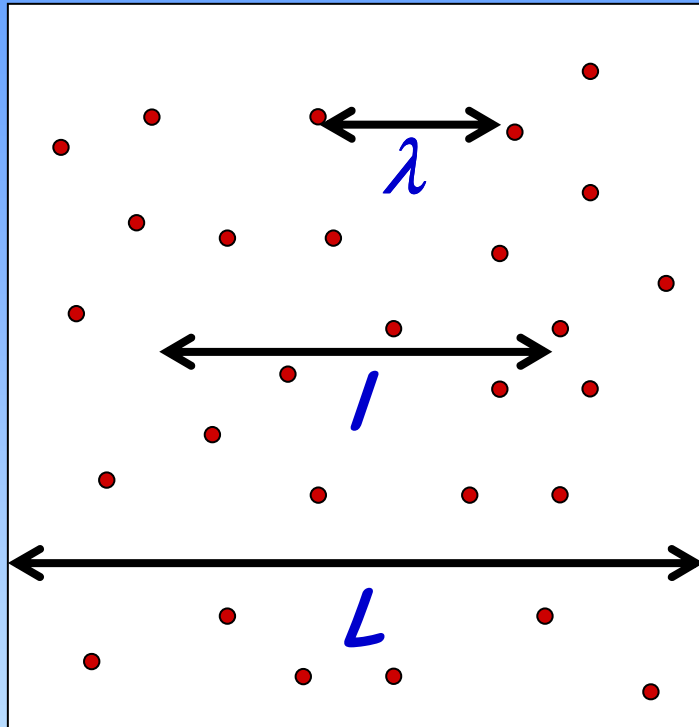
drift of cluster towards wall !

In an isolated 1D case  
granular hydrodynamics  
does not work.

What about the general case?



# Knudsen number



$\lambda$  = mean free path

$l$  = typical length at which  
macroscopic quantities vary

$L$  = typical system size

$Kn = \lambda/L$  (global Knudsen  
number)

$Kn_{loc} = \lambda/l$  (local Knudsen  
number)

**Hydrodynamics work if  $Kn \ll 1$  !**

Molecular system: local  $Kn \ll 1$  (not a Knudsen gas!)

Granular system: local  $Kn$  large !

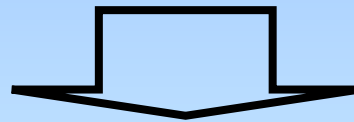
# No separation of scales

No separation of length scales:

macroscopic quantities vary on the same scale  
as the mean free path !

Flowing systems:

mean velocity  $\sim$  "thermal" velocities = velocity  
fluctuations



No separation of time scales:

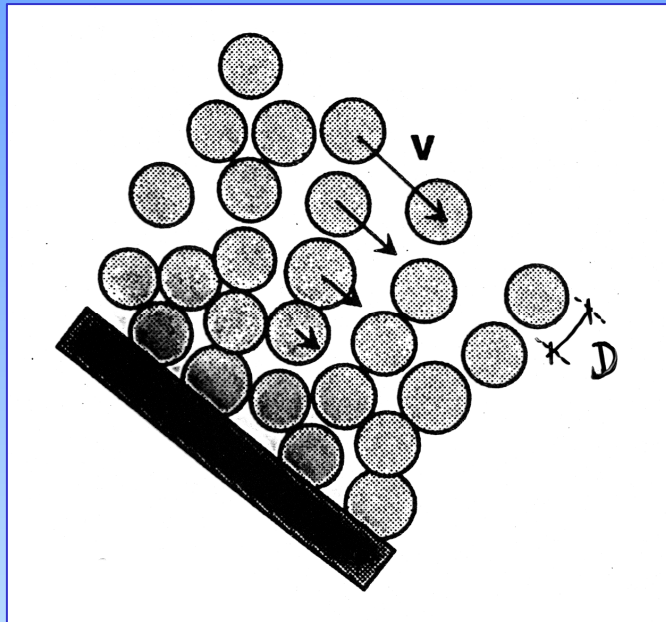
macroscopic quantities can change as fast as  
particle velocities

# "Conclusion"

There are many reasons why granular hydrodynamics should NOT work...

The surprise is that nevertheless in many cases it DOES work !!!!

# Influence of the Bagnold regime



$$\gamma = \partial_z v =$$

$D\gamma = \text{typ}$   
respect

$\lambda_e = \text{cha}$   
over wh  
due to f



*Ralph Bagnold*

Friction force  $F_c$ :

$$F_c \approx \frac{\Delta \epsilon_k}{\lambda_e} \sim \frac{m(D\gamma)}{\lambda_e}$$

Viscous damping force  $F_v$ :

$$F_v = 3\pi\eta Dv \sim \eta\gamma D^2$$

Bagnold number:

$$B = \frac{F_c}{F_v} \sim \frac{m\gamma}{\lambda_e\eta}$$

$B > 450$ : granular regime

$B < 10$ : ambient fluid crucial

# Bagnold number for a vertically vibrated granular gas

For a vertically vibrated granular gas the Bagnold number is defined as the ratio of gravity and viscous forces:

$$B = \frac{F_z}{F_c} = \frac{mg}{3\pi\eta Dv} = \frac{\rho_g D^2 g}{18\eta v}$$

For a sand/glass particle ( $\rho = 2.5 \cdot 10^3 \text{ kg m}^{-3}$ ) in air ( $\eta = 1.9 \cdot 10^{-5} \text{ Pa s}$ ) with a typical velocity of 1 m/s we have:

$$\left. \begin{array}{l} B \approx 7.3 \cdot 10^7 D^2 \\ B \approx 1 \end{array} \right\} \rightarrow D \approx 100 \mu\text{m}$$



# Sand dunes: "barchans"

Marokko



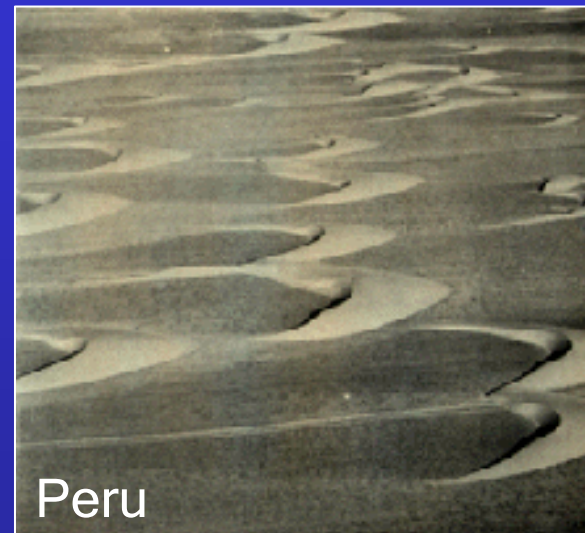
Tunesiä



Namibiä



Peru





# Sand dunes travel ...

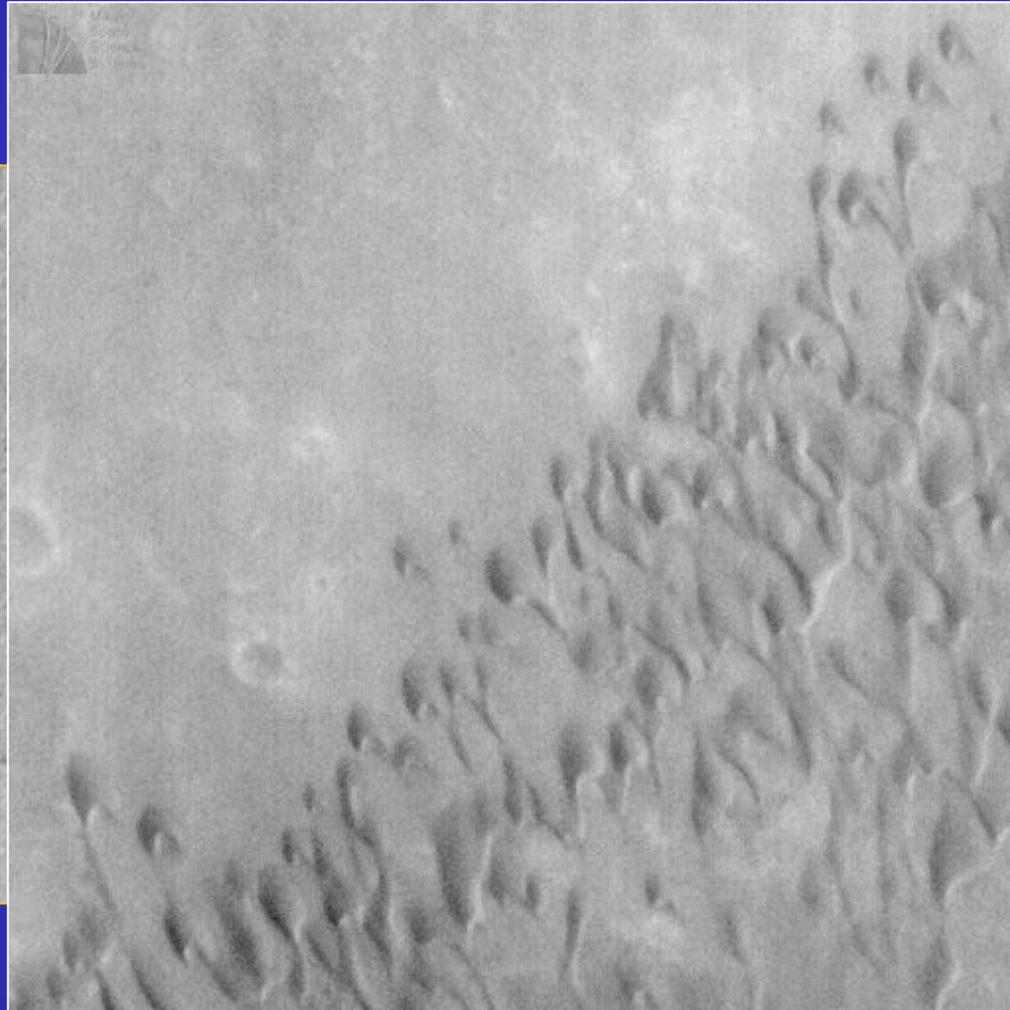


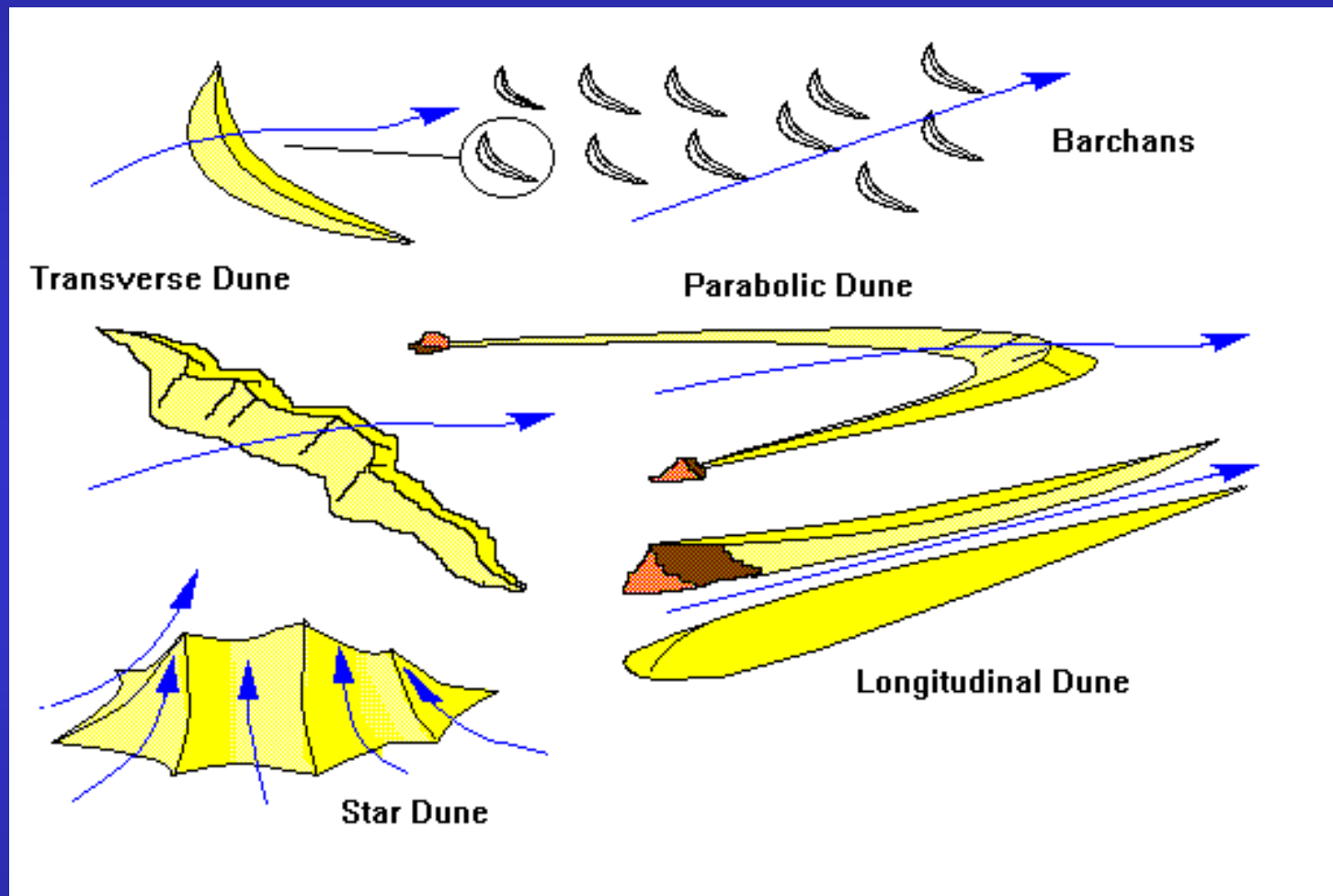
Egypt

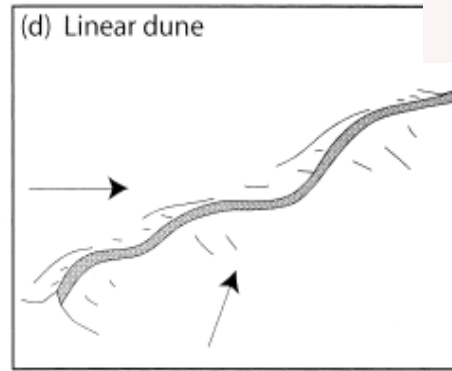
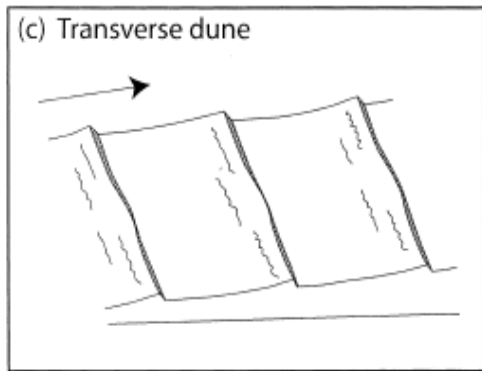
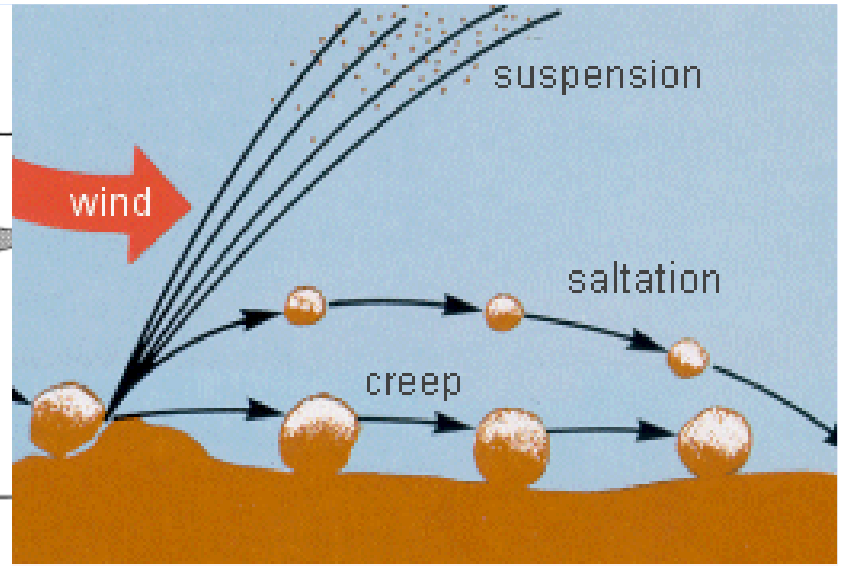
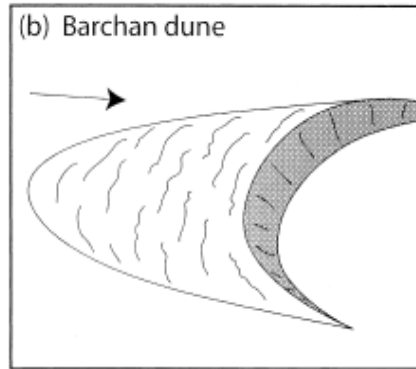
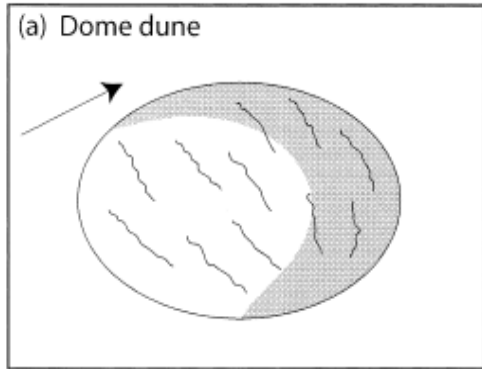
# Barchans op Mars



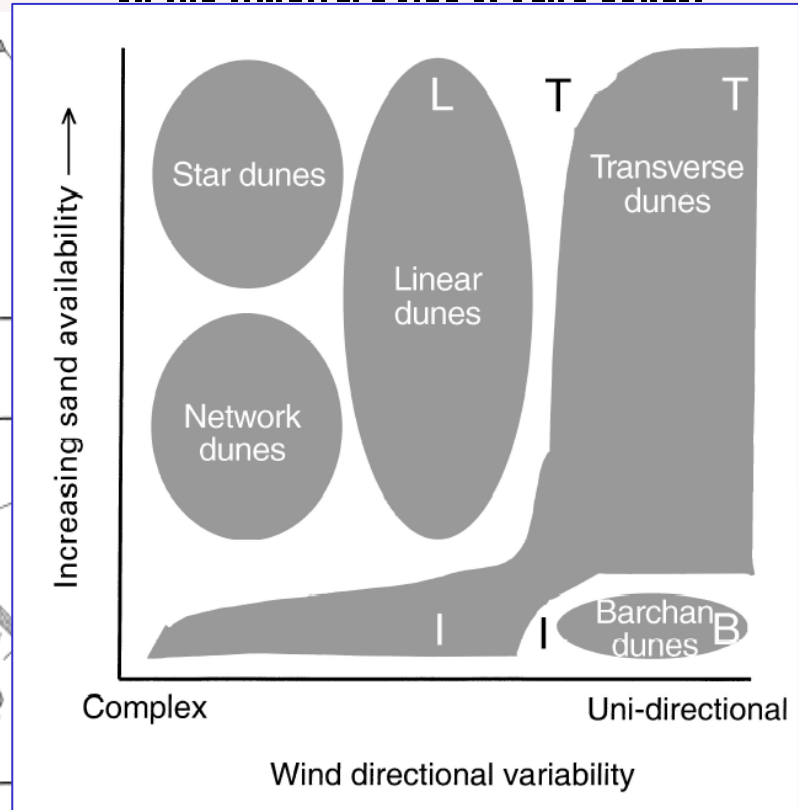
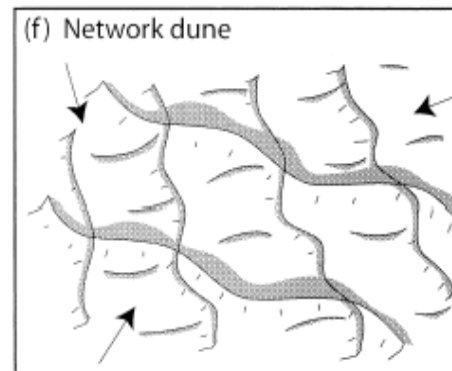
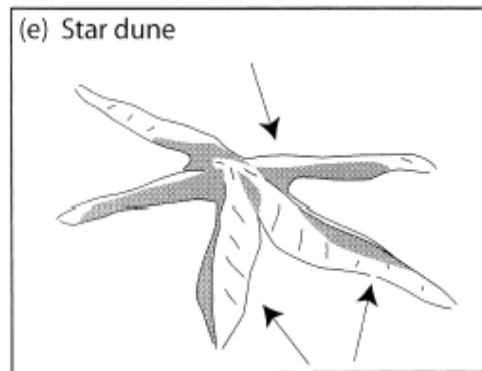
Photos from the Mars Orbiter Camera  
(Mars Global Surveyor project)



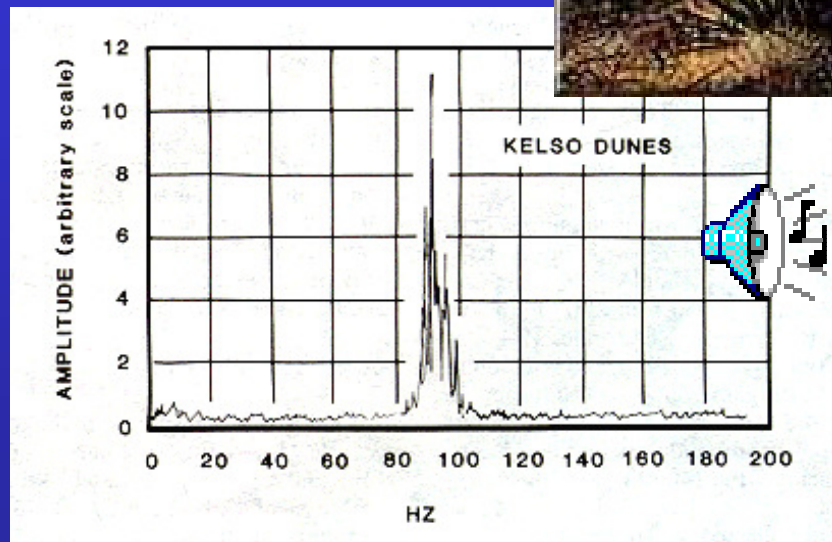




**Wind causes saltation, or jumping grains, on the windward side of sand dunes.**

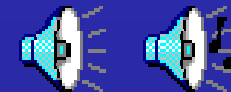


# Singing sand

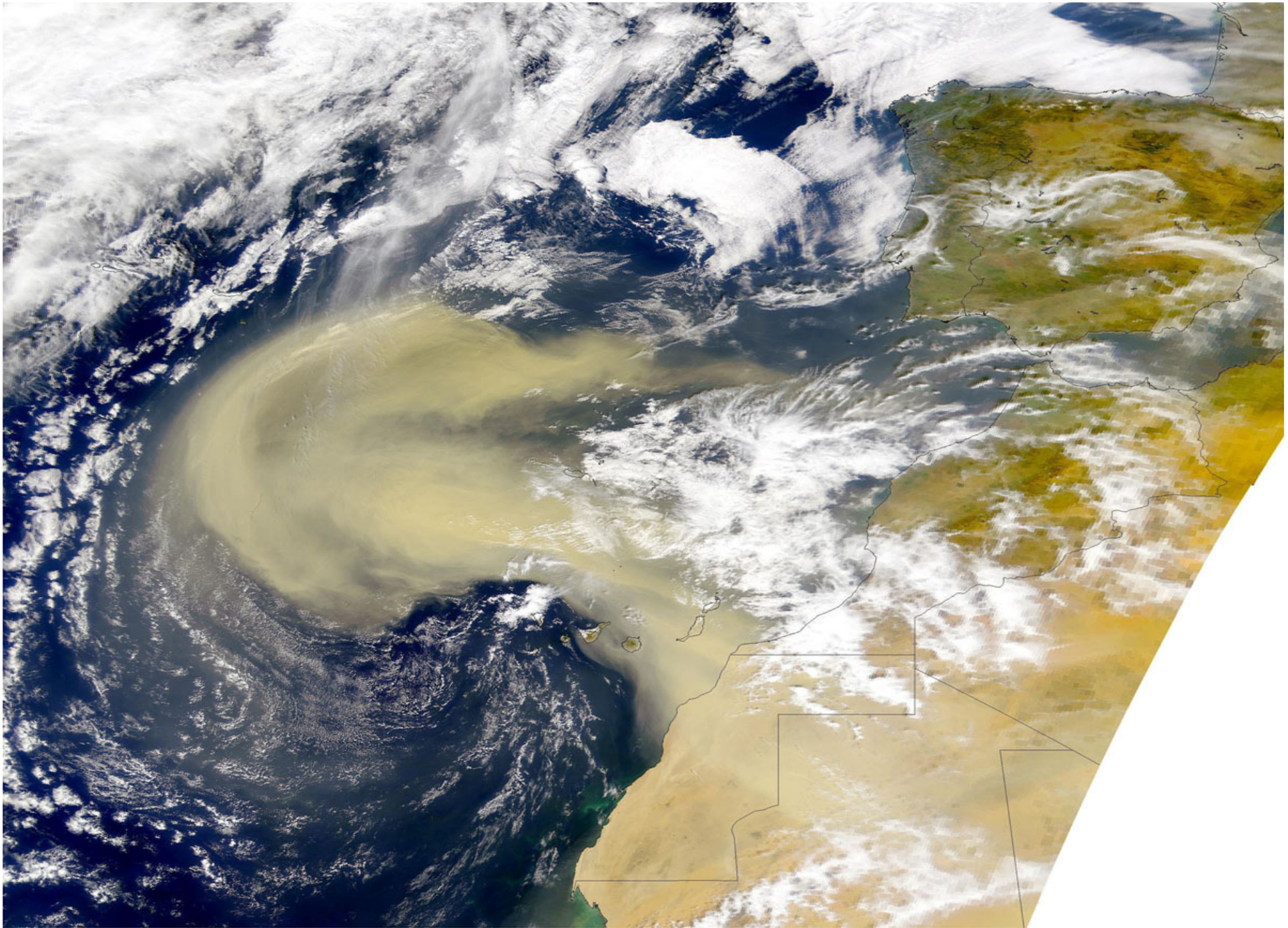


← Frequency spectrum of "booming sand" in the Kelso dunes, California.

Sand can also chirp or quack →







**SeaWiFS Captures Massive Dust Storm**

**February 26, 2000**

**SeaWiFS Project/ORBIMAGE**



# References

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